

Neogene Uplift of The Barents Sea

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A. Amantov*

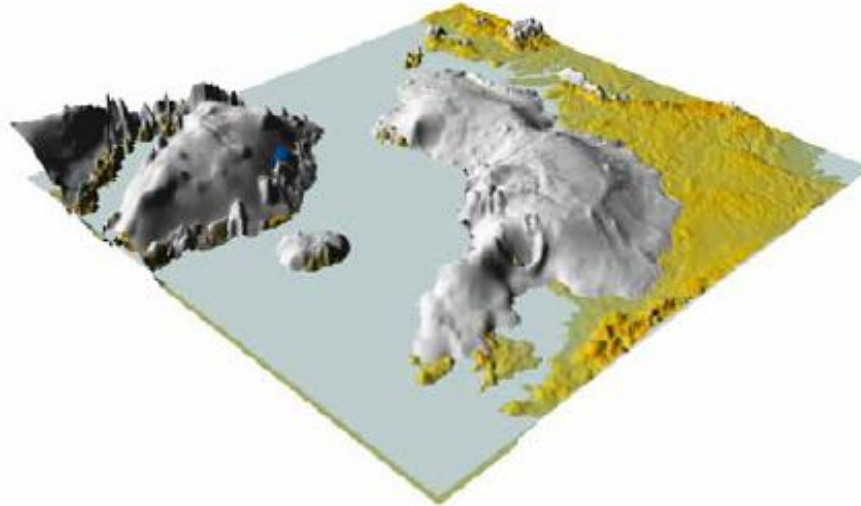
Tectonor/UiS, Stavanger, Norway

FORCE seminar April 4, 2013

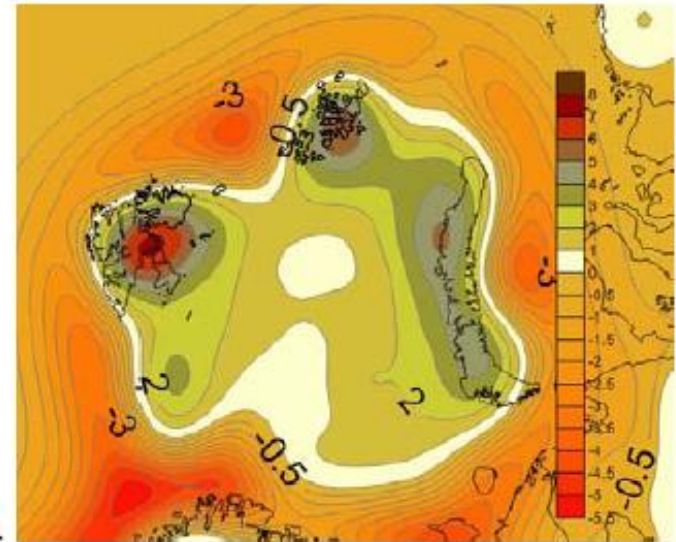


The project (2010-2012)

Neogene Uplift of the Barents Sea



a.



b.

Figure 1a. Modelled ice sheet at Last Glacial Maximum; 1b. Calculated present rate of uplift for the Barents Sea (mm/yr).

**A study by Tectonor, IRIS, VSEGEI, Royal Holloway University of London,
and Cornell University**

Funding companies



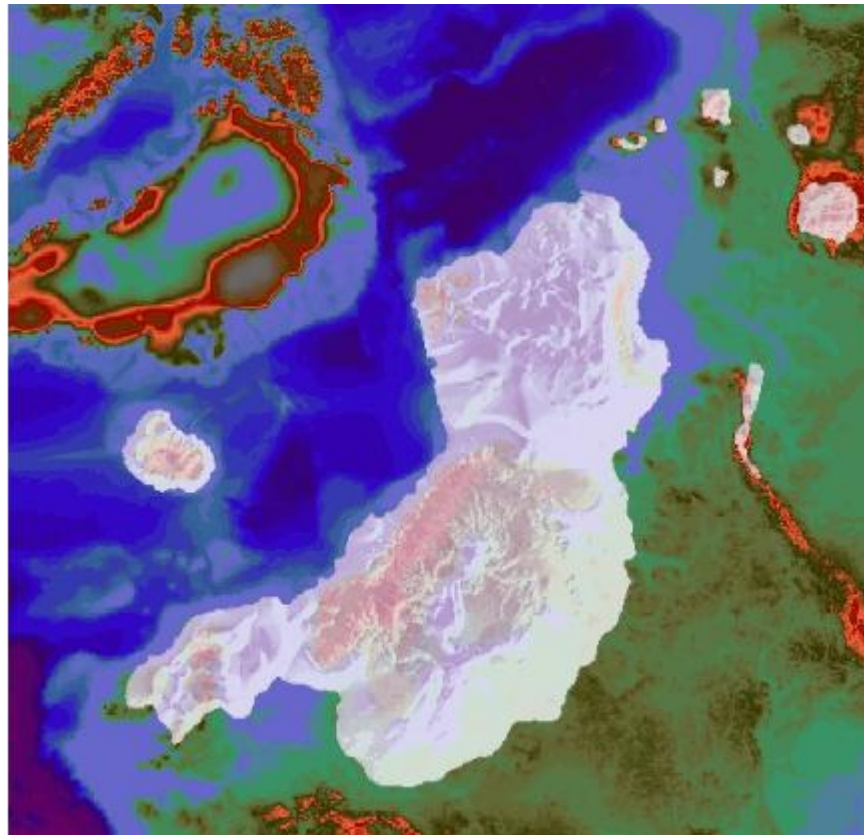
Objective

The objective of the work was to improve our understanding of the timing, geometry and magnitude of the Neogene and Quaternary erosion and uplift of the Russian and Norwegian parts of the Barents Sea.

- ✓ **Plio-Pleistocene glaciations - extent and timing**
 - ✓ **Effects of glaciers**
- ✓ **Estimates/modelling of glacial erosion**
 - ✓ **Isostatic effects (tilting of reservoirs)**
- ✓ **Temperature- og maturity effects**
- ✓ **Estimations of Neogene Erosion**

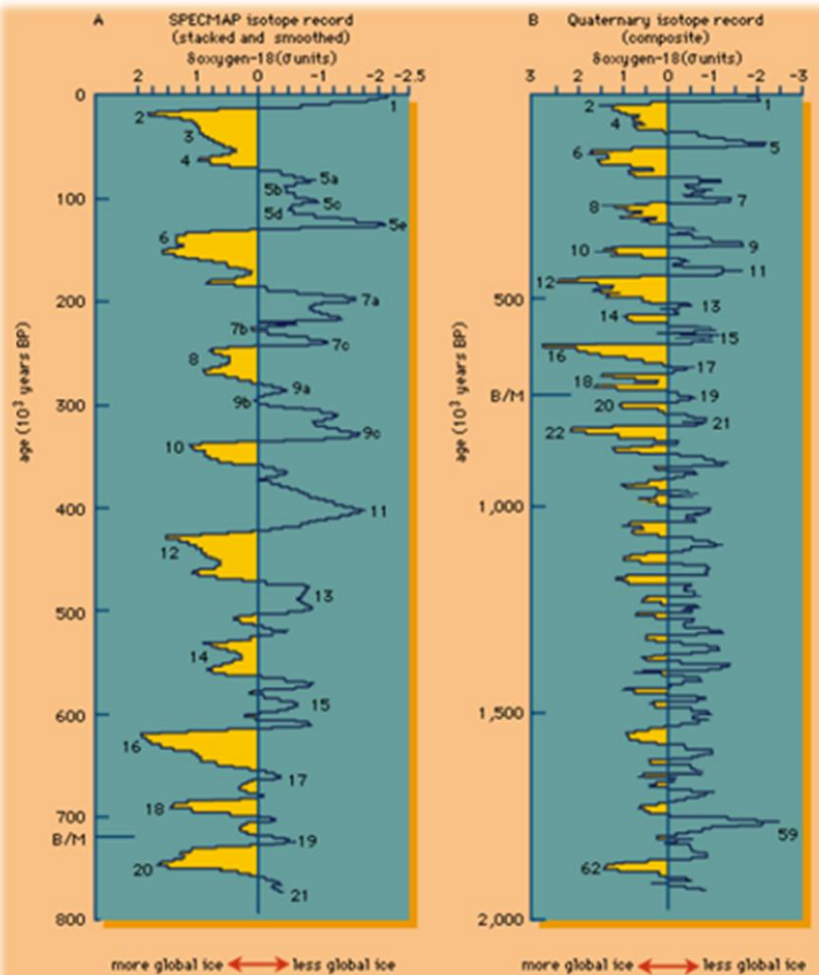
Plio-Pleistocene glaciations

Deglaciation - last glaciation



20 000 BP

Glacial-Interglacial cycles



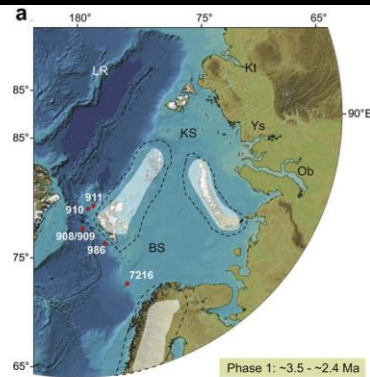
(A) The SPECMAP (Spectral Mapping Project) record based on five low- and middle-latitude deep-sea cores and (B) a composite record of four cores from the equatorial Pacific, the Caribbean, and the North Atlantic. Isotopic stages and substages are indicated; B/M shows the level of Brunhes/Matuyama reversal.

From deep sea sediments

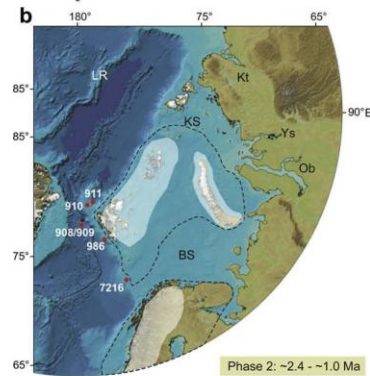
At least 8 glacial-interglacial cycles over the last 800 000 years, and maybe 30 cycles since late Pliocene.

3 Plio-Pleistocene phases

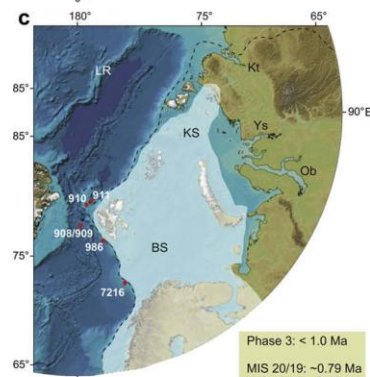
Typical glaciations



3.5-2.4 Ma



2.4-1.0 Ma

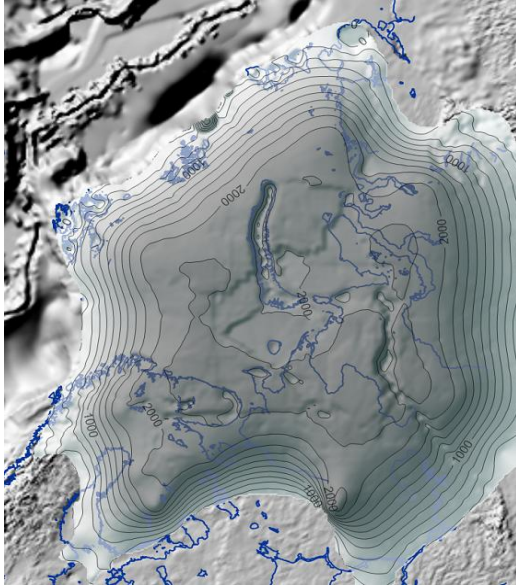


< 1.0 Ma

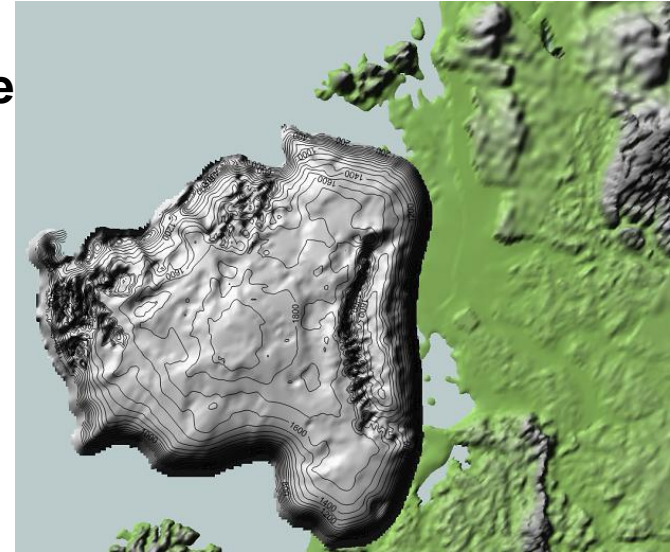
Knies et al (2009)

Plio-Pleistocene Glaciations

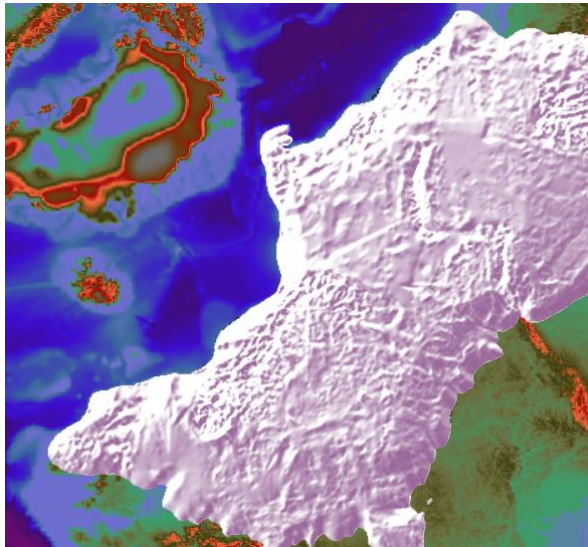
1st phase



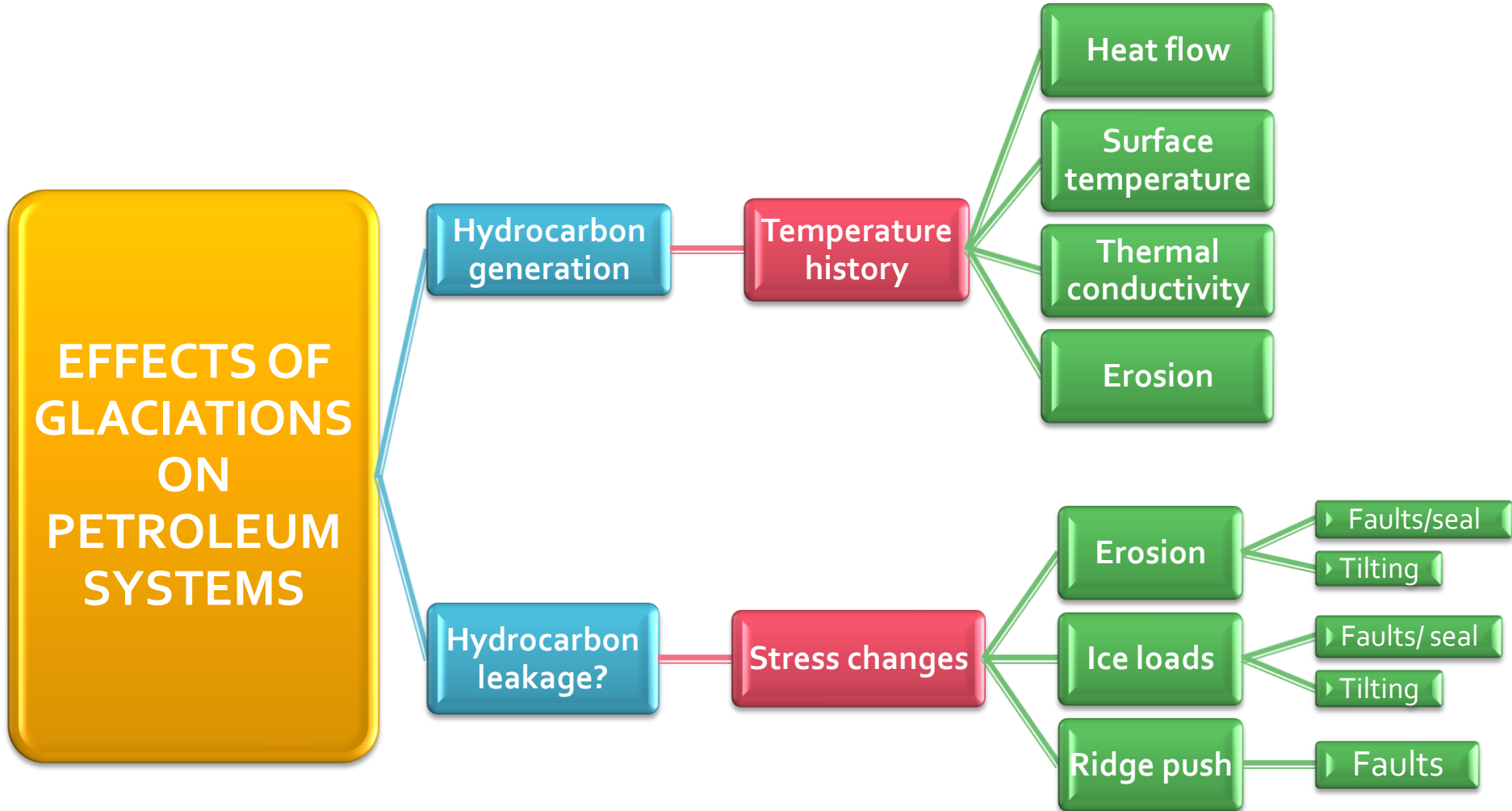
2nd phase



3rd phase

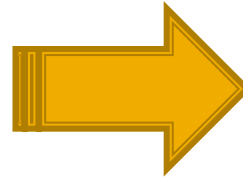


Maximum possible ice extent and thicknesses for three periods of Plio-Pleistocene



Glaciations

- ✓ Surface temperature
- ✓ Thermal conductivity
- ✓ Compaction
- ✓ Isostatic effects
- ✓ Elastic effects
- ✓ Stress effects
- ✓ Erosion

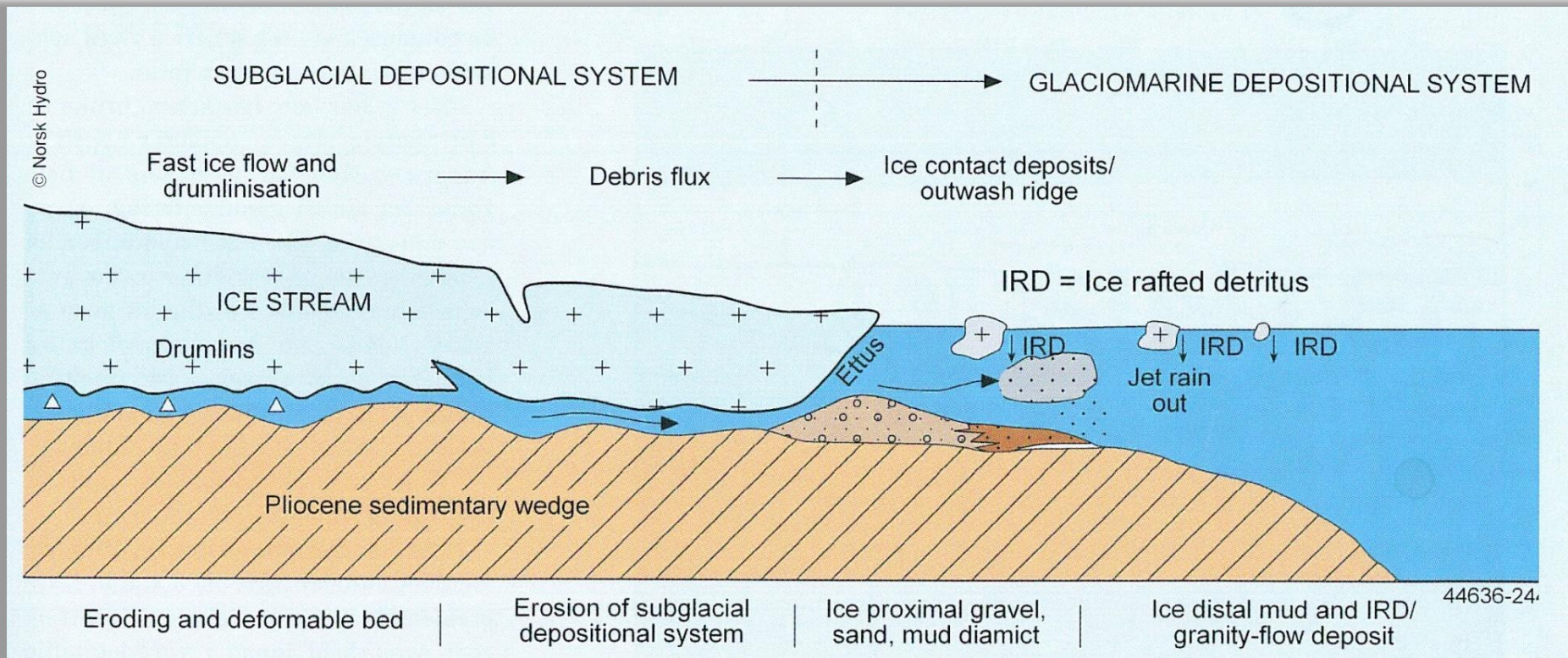


Petroleum systems

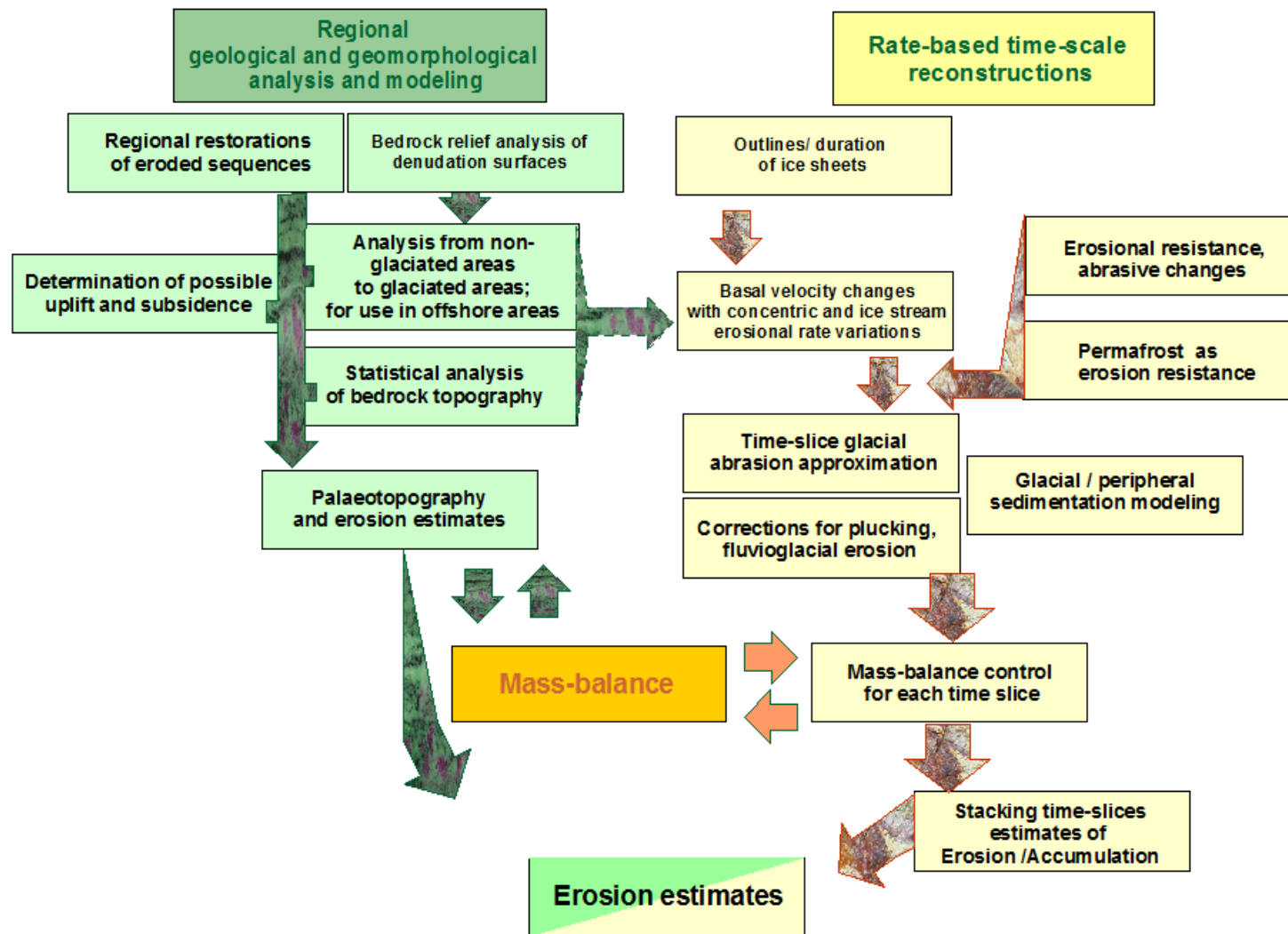
- ✓ temperature
- ✓ maturity
- ✓ migration routes
- ✓ leakage
- ✓ reservoir quality

Glacial erosion

Effects of glaciations – Glacial erosion

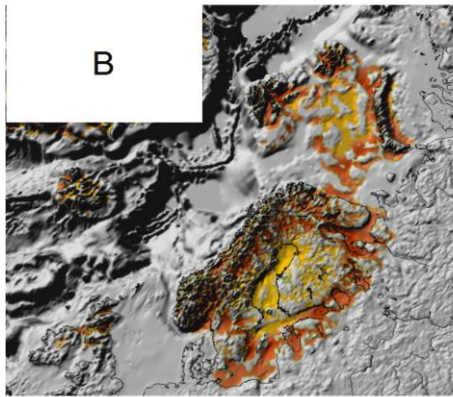
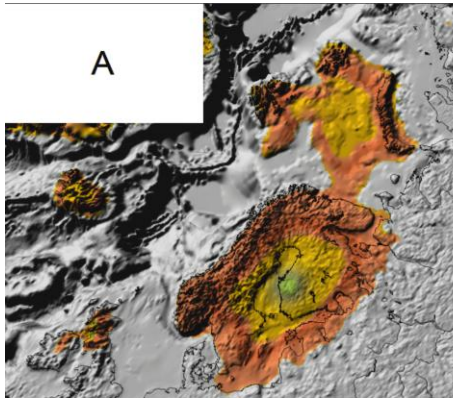


Morphological modelling

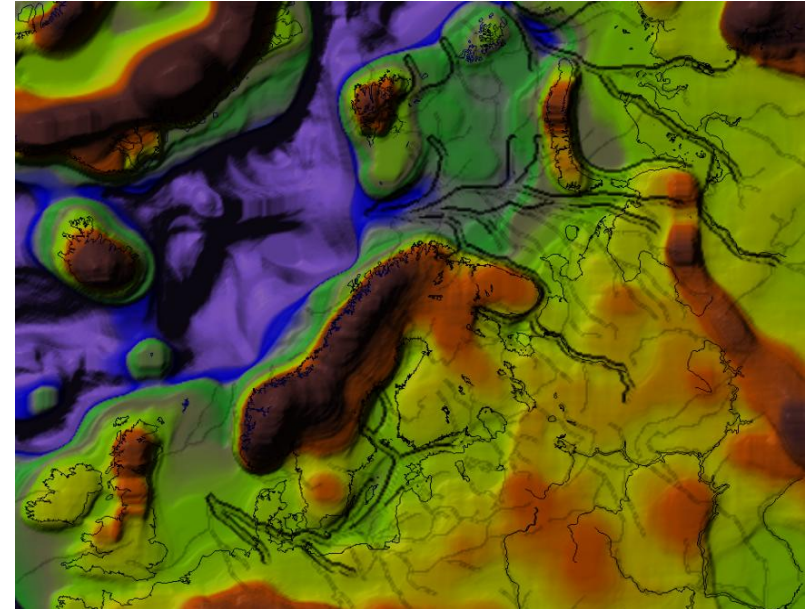


Morphological modelling

Concentric pattern forms due to low ice velocity under the center, and more rapid basal ice velocity near the margins



Effect of ice streams with enhanced erosional capacity



Model of pre-glacial landscape with major drainage pattern

Different erodability of sedimentary bedrock and basement lithologies

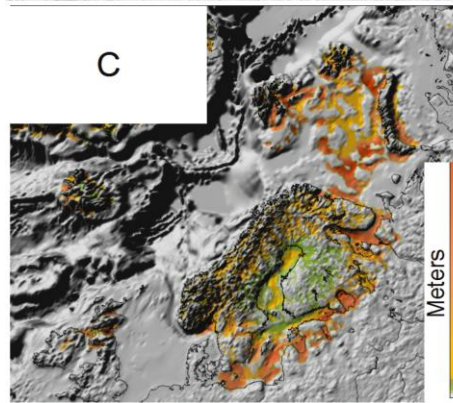
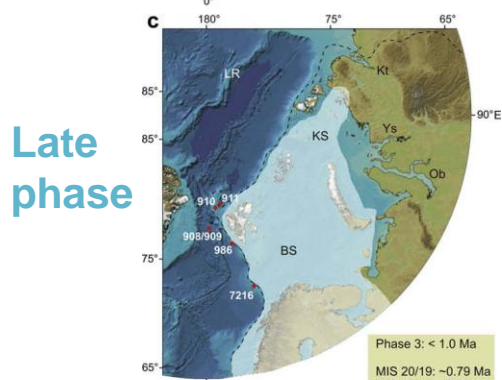
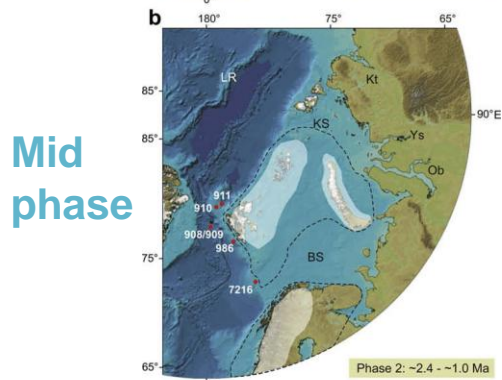
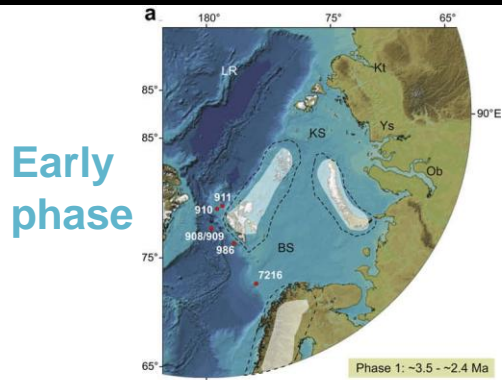


Illustration of the erosion model

Erosion modeling

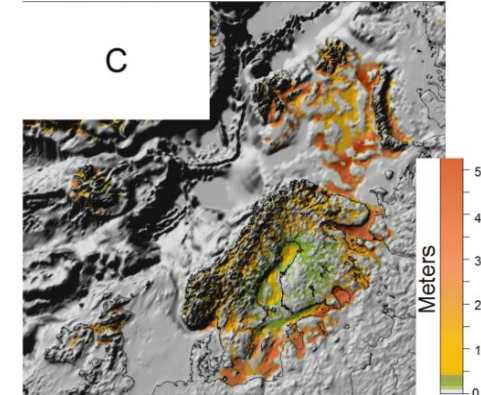
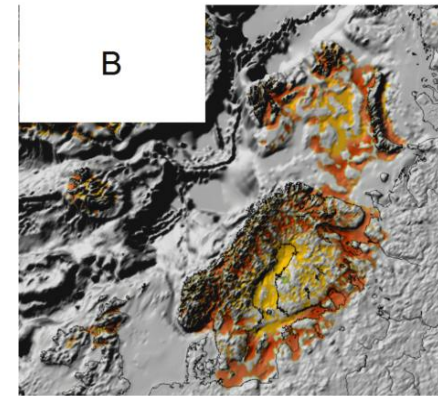
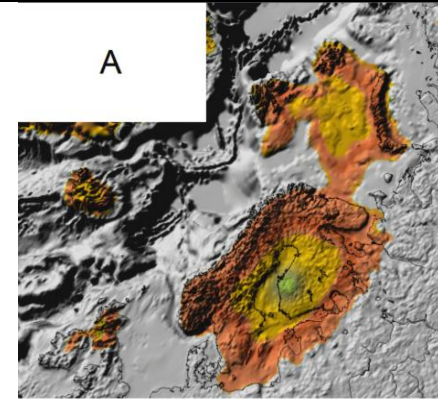


Knies et al (2009)

- MIS 20/19: -0.79 Ma
- MIS 16/15: -0.62 Ma
- MIS 12/11: -0.41 Ma
- MIS 6/5: -0.13 Ma
- MIS 5.4: -0.1 Ma
- MIS 2/1: -0.015 Ma

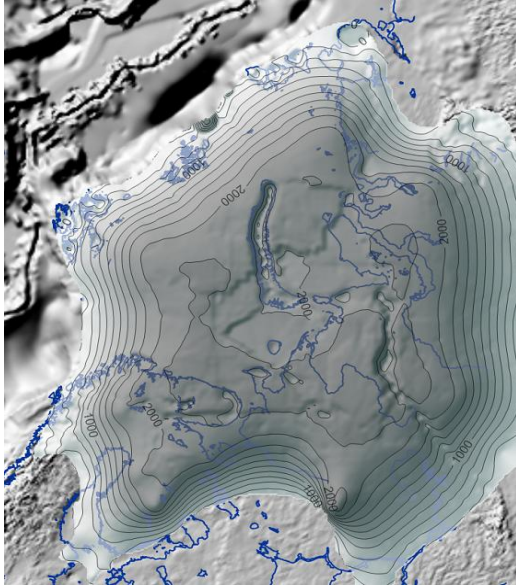


Predict erosion

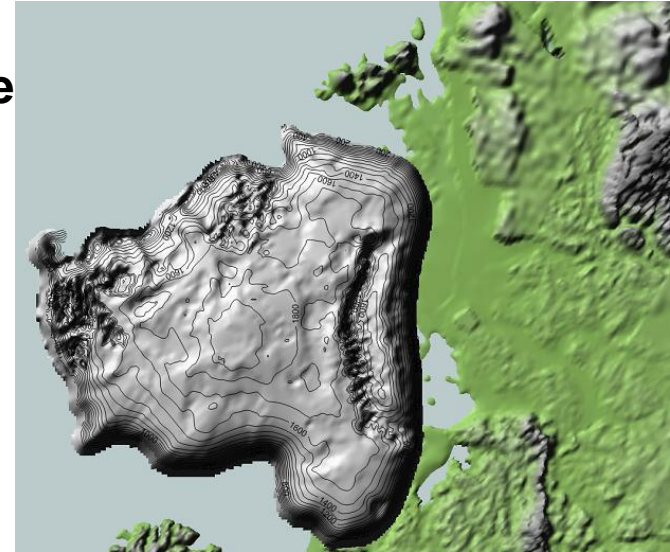


Plio-Pleistocene Glaciations

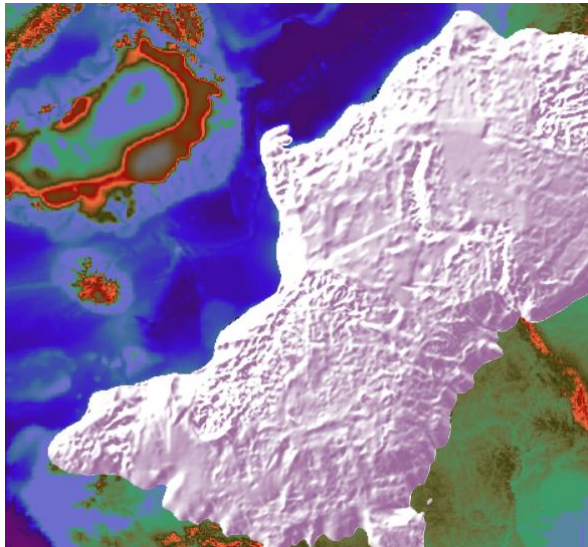
1st phase



2nd phase

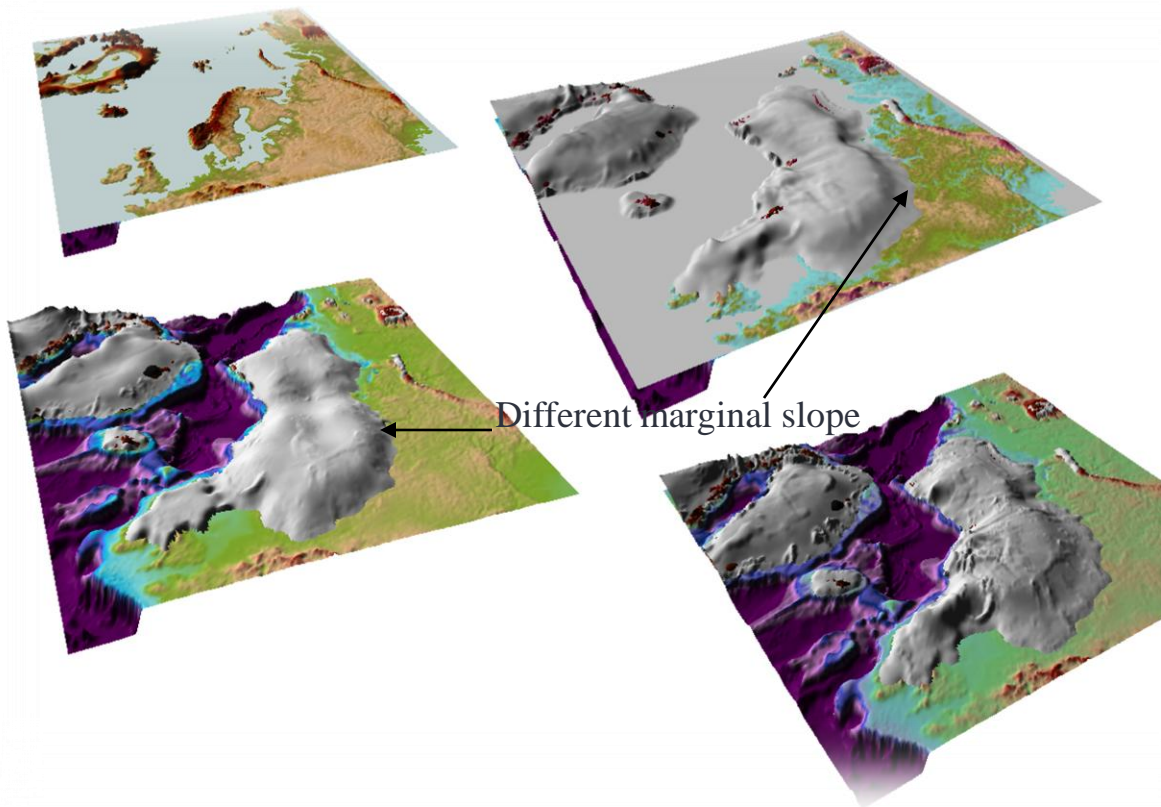


3rd phase



Maximum possible ice extent and thicknesses for three periods of Plio-Pleistocene

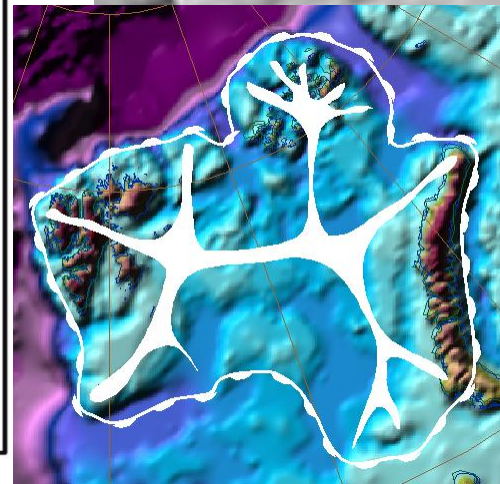
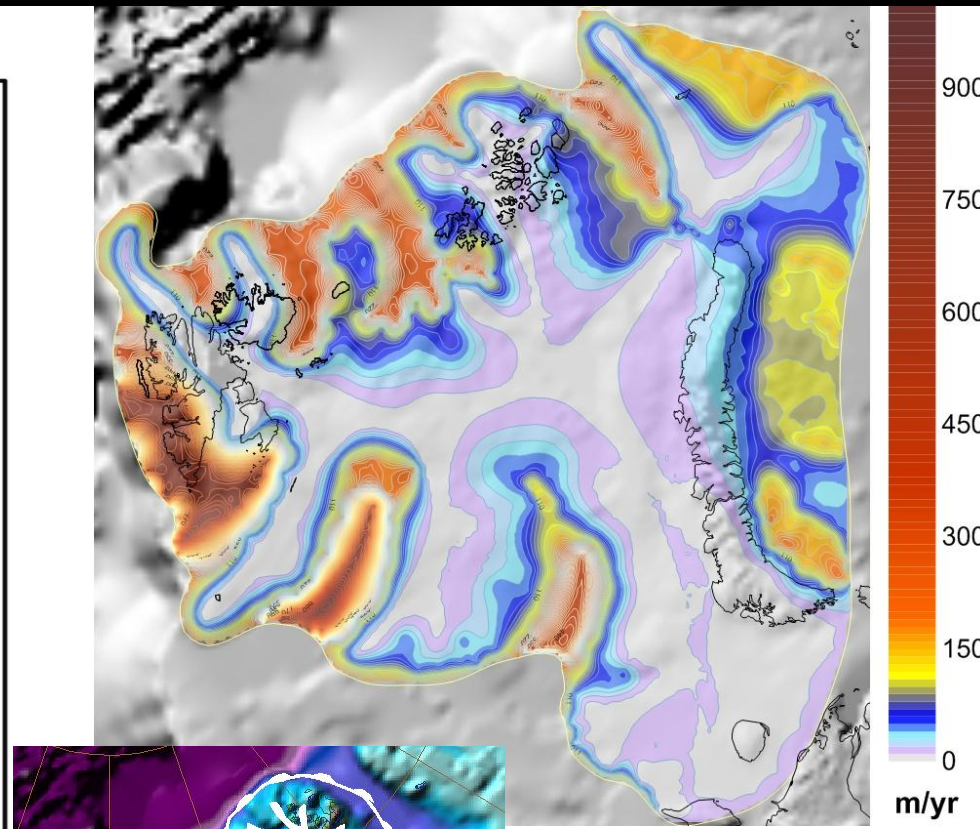
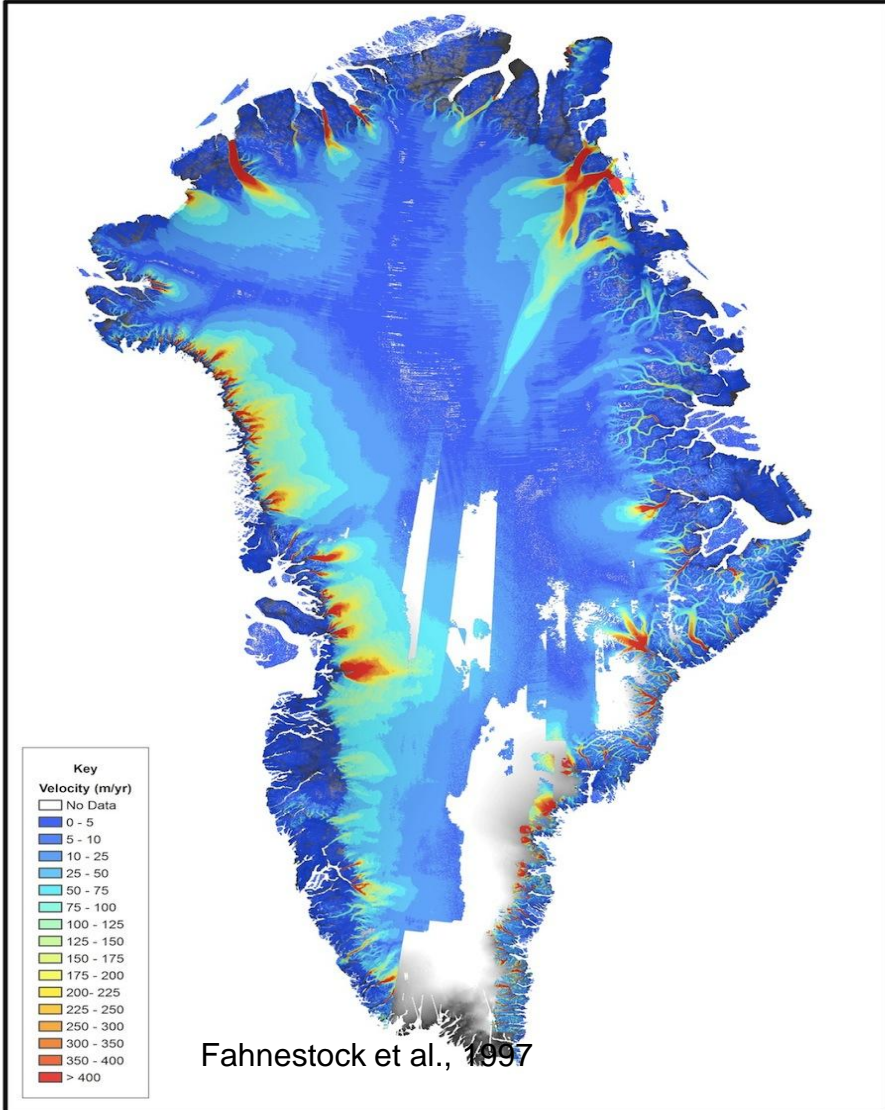
Ice Thickness Module



Ice Thickness computed from:

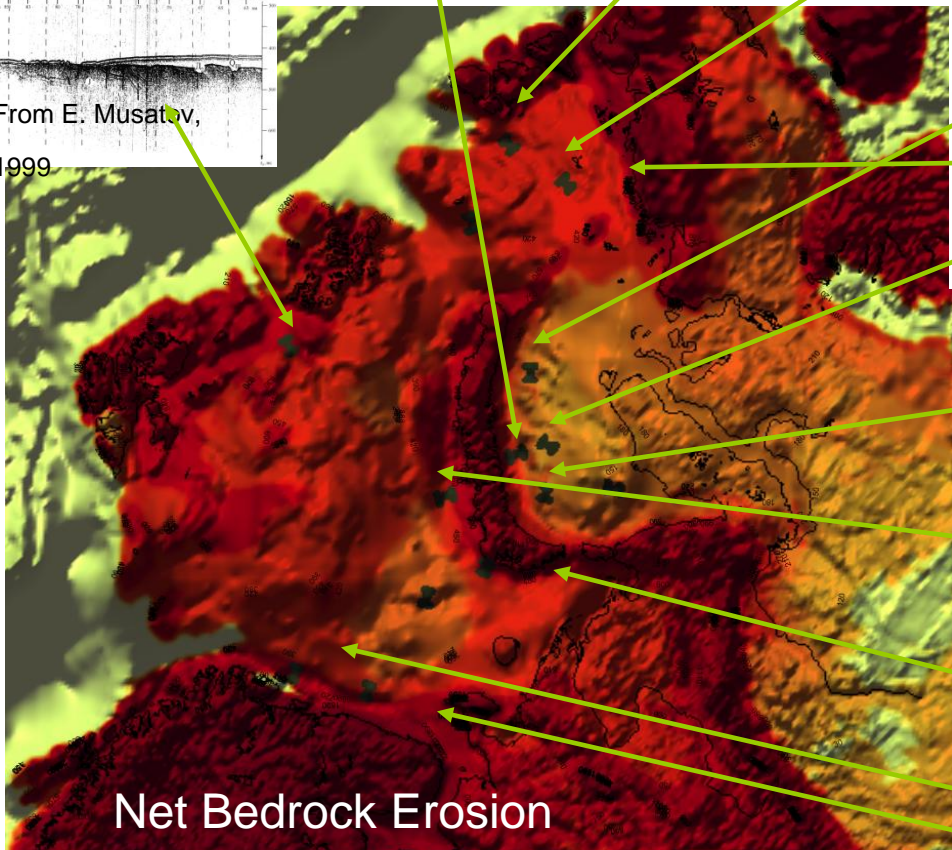
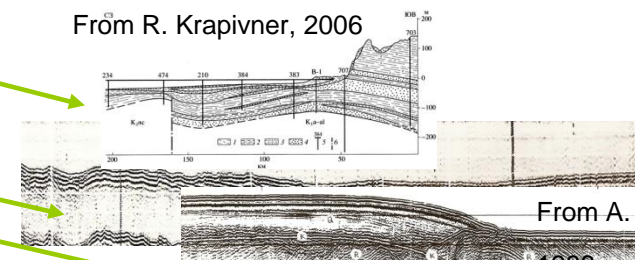
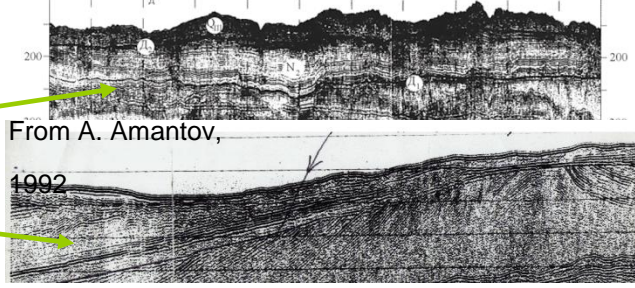
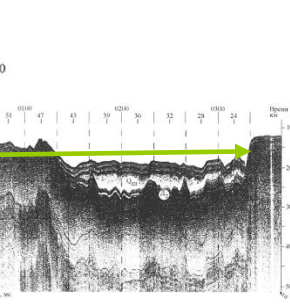
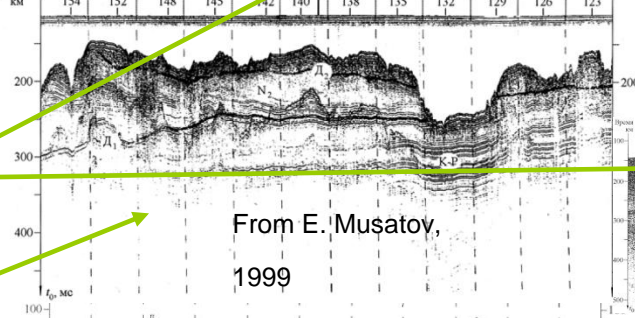
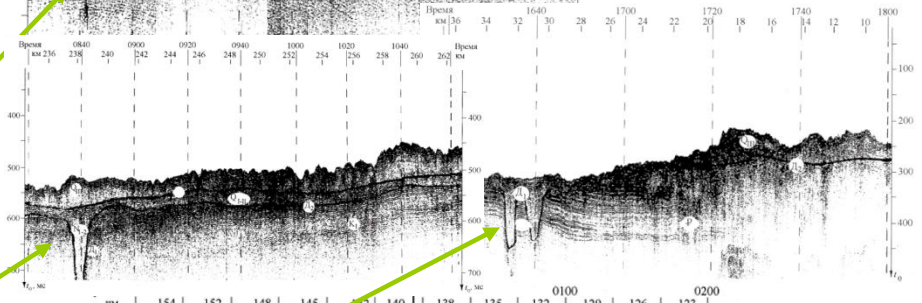
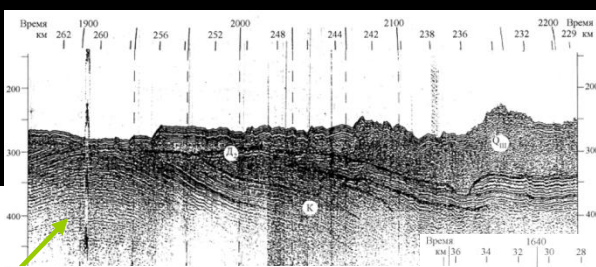
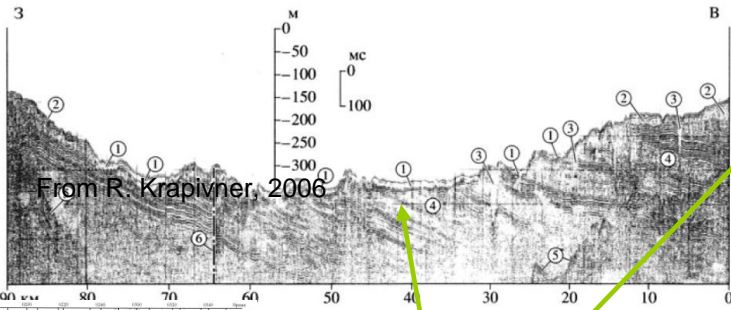
1. Ice margin outline
2. Topography
3. Basal lithology
4. Ice velocity
5. Floating or frozen base
6. Marginal slope (specified)
7. Continental or ocean margin

Ice surface velocities



Skeleton

Shallow seismic

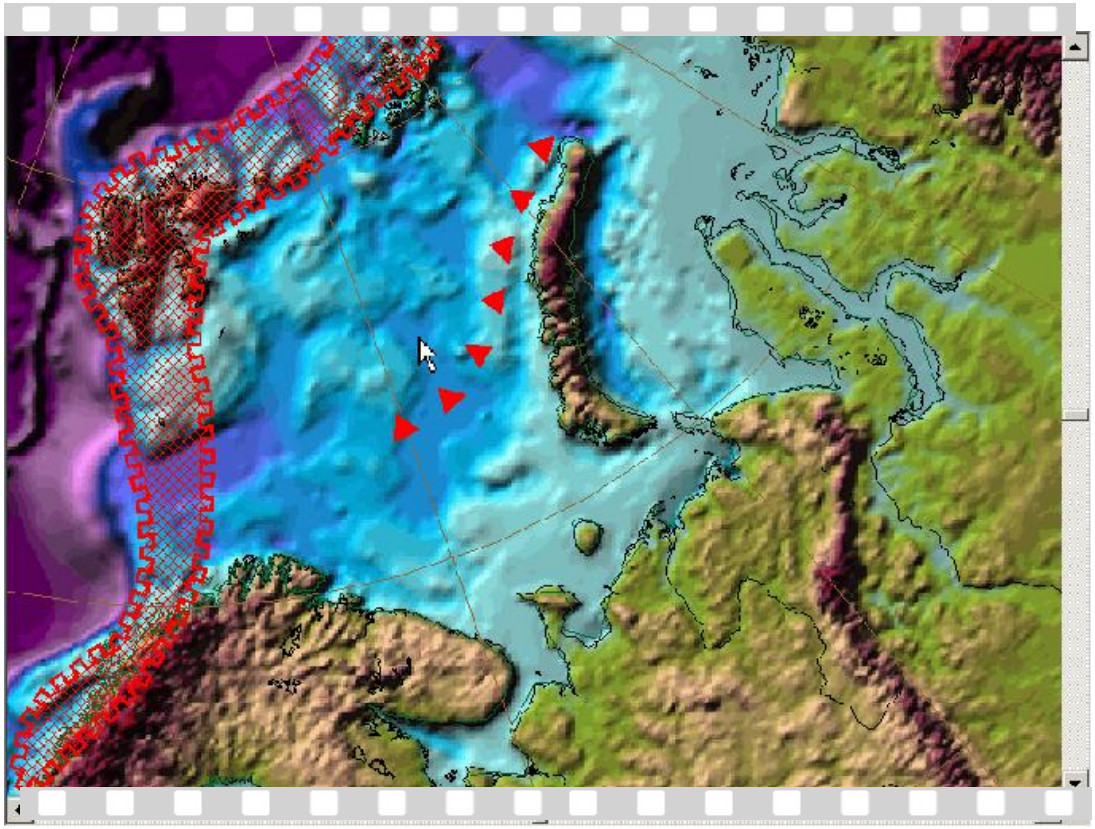


Net Bedrock Erosion

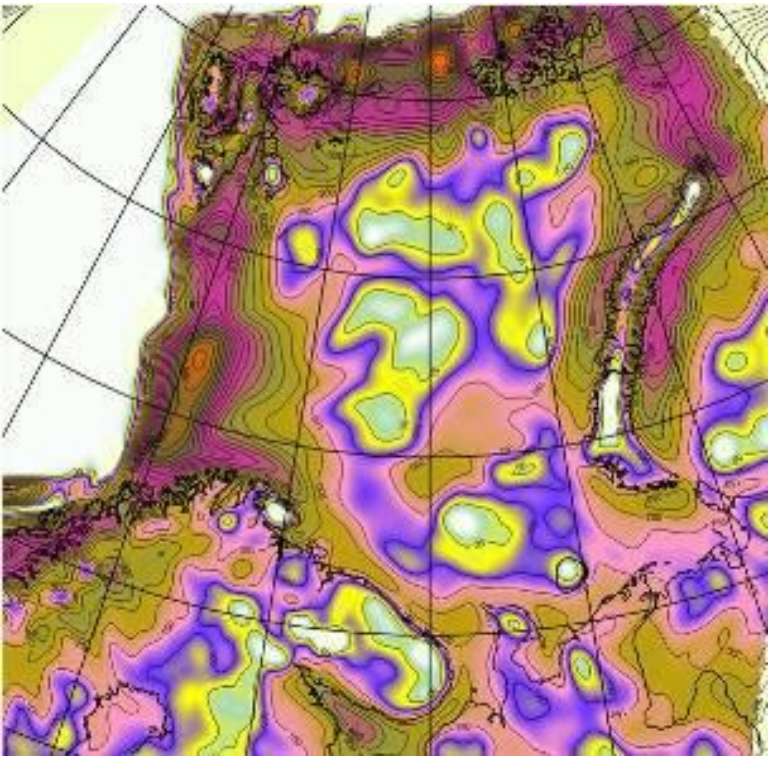
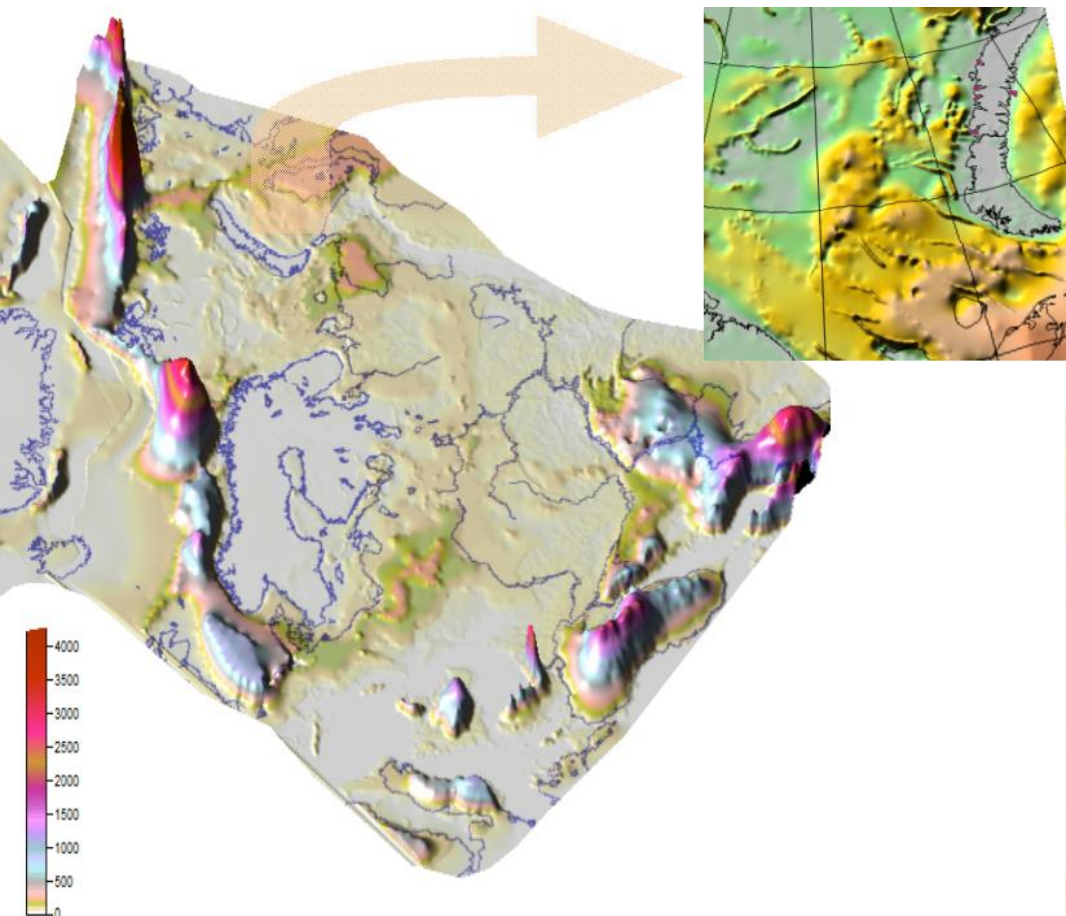


From A. Amantov, 1988

Zone of maximum glacial erosion



Maximum glacial erosion has stable position along the Atlantic coast, while eastern flank is migrating due to ice sheet grow and decay



Plio-Pleistocene
erosion and deposition
(Amantov, unpublished)

Isostatic response

Archimedes of Syracuse

(Greek: [Ἀρχιμήδης](#))

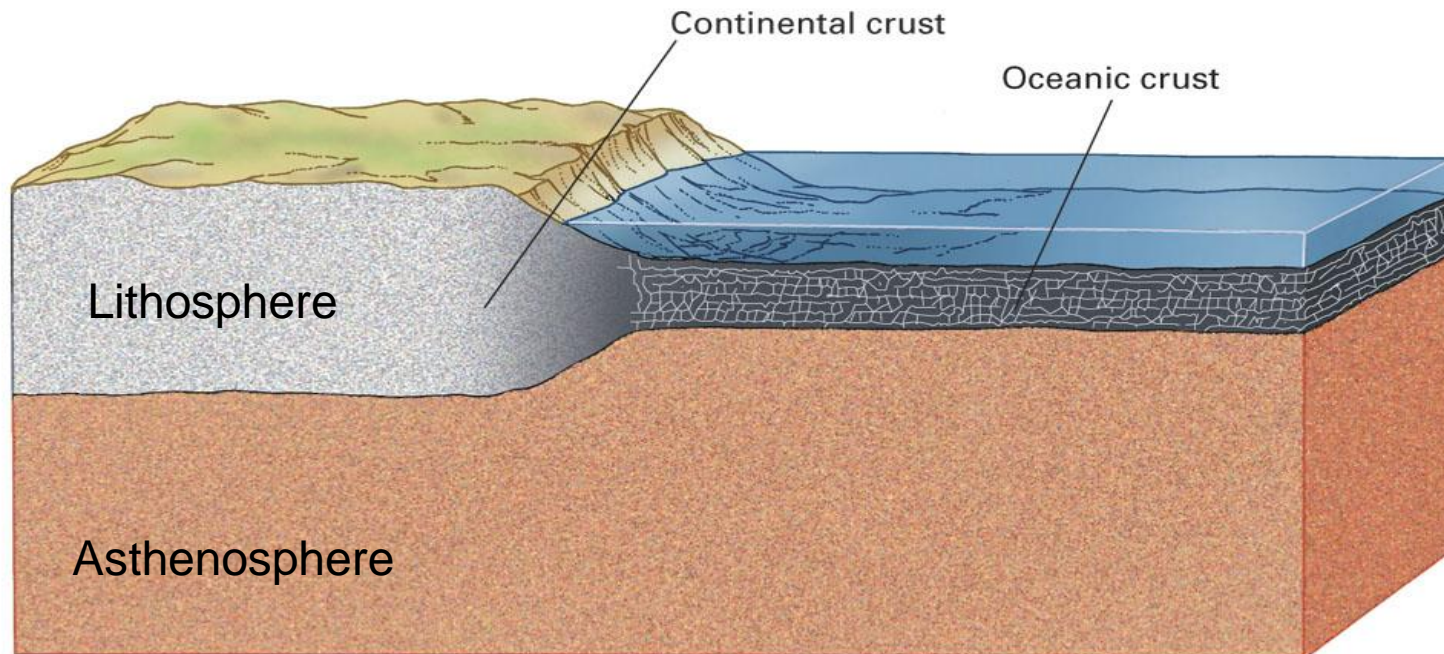
Archimedes is generally considered to be the greatest mathematician of antiquity and one of the greatest of all time



A buoyancy force arises when a solid object is placed into an (ideal)-liquid.

The buoyancy force is specified by Archimede's Principle which states: the decrease in weight of the object equals the weight of the liquid displaced by the submerged portion of the body.

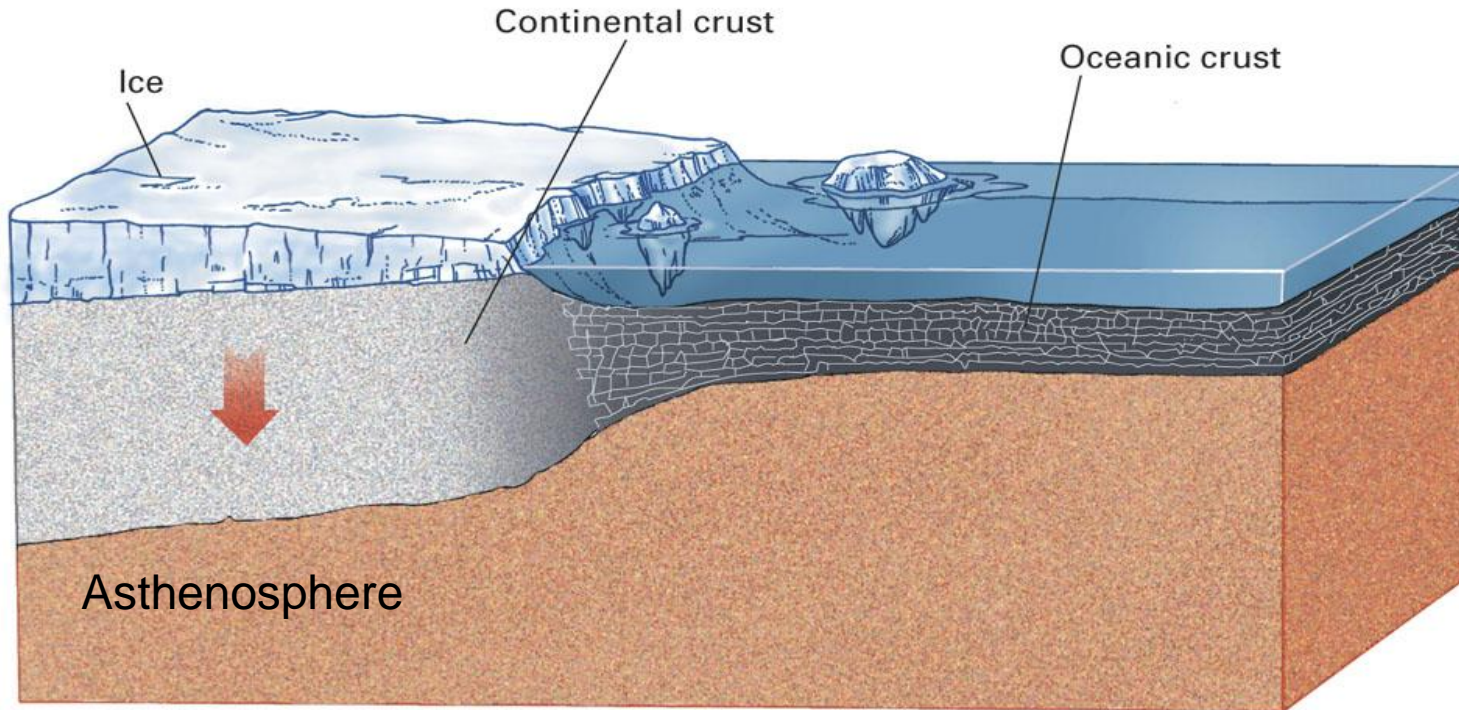
Isostasy



© 2005 Brooks/Cole - Thomson

- The lithosphere in floating equilibrium on the asthenosphere is isostasy

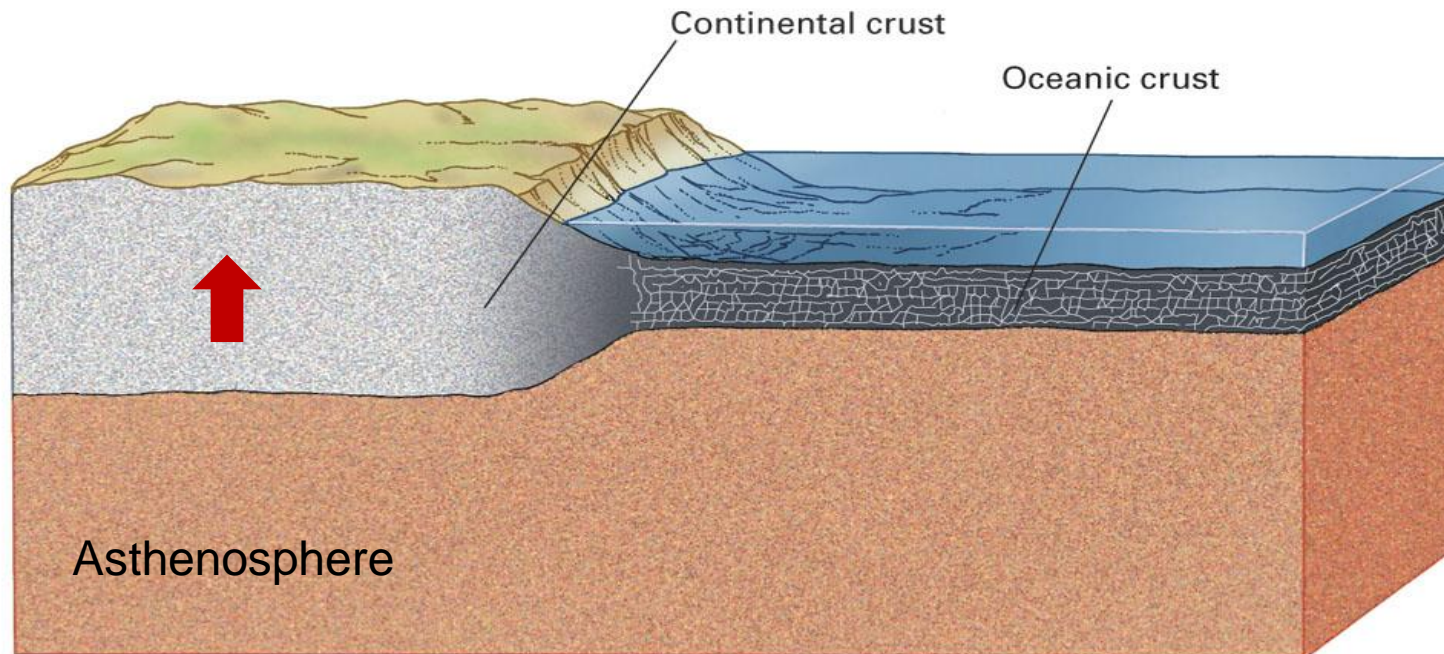
Glacial Isostatic Adjustment



© 2005 Brooks/Cole - Thomson

- Vertical movement in response to changing burden is called isostatic adjustment
- The crust is subsiding due to the ice load

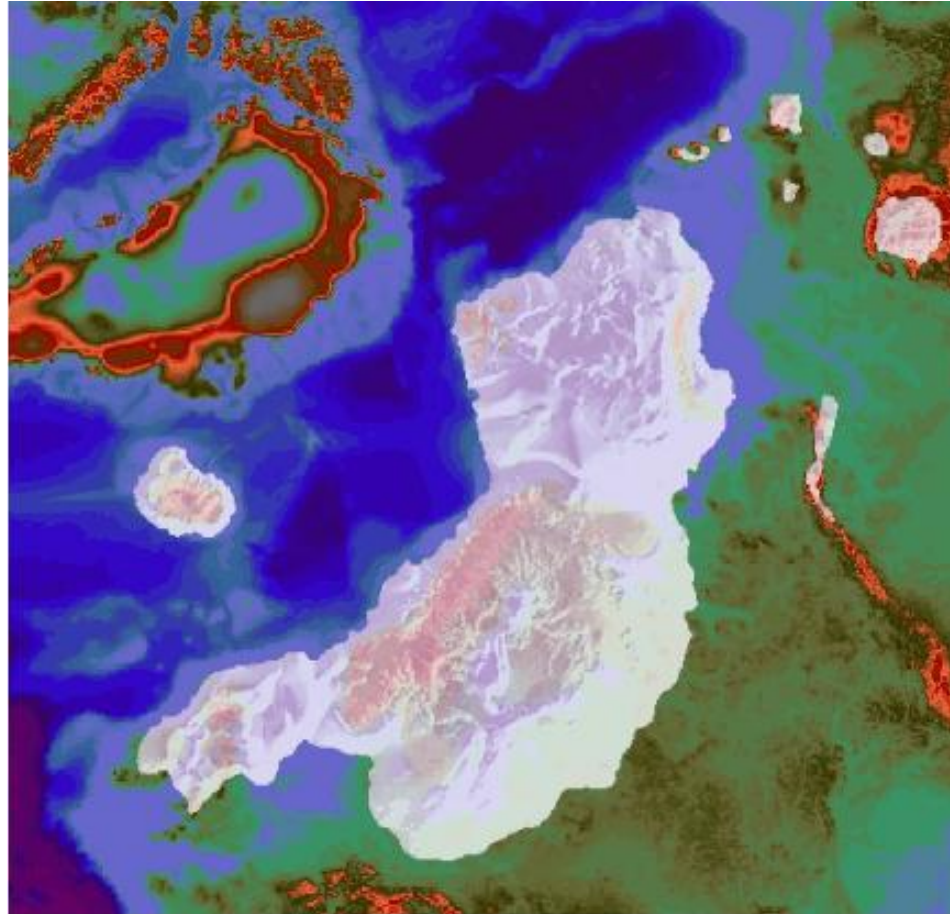
Glacial rebound



© 2005 Brooks/Cole - Thomson

- When the load is removed, the crust is experiencing uplift - until new equilibrium is established

Deglaciation model ('AA1')



20 000 BP

Calculated glacial isostasy

last 20 000 years



19000

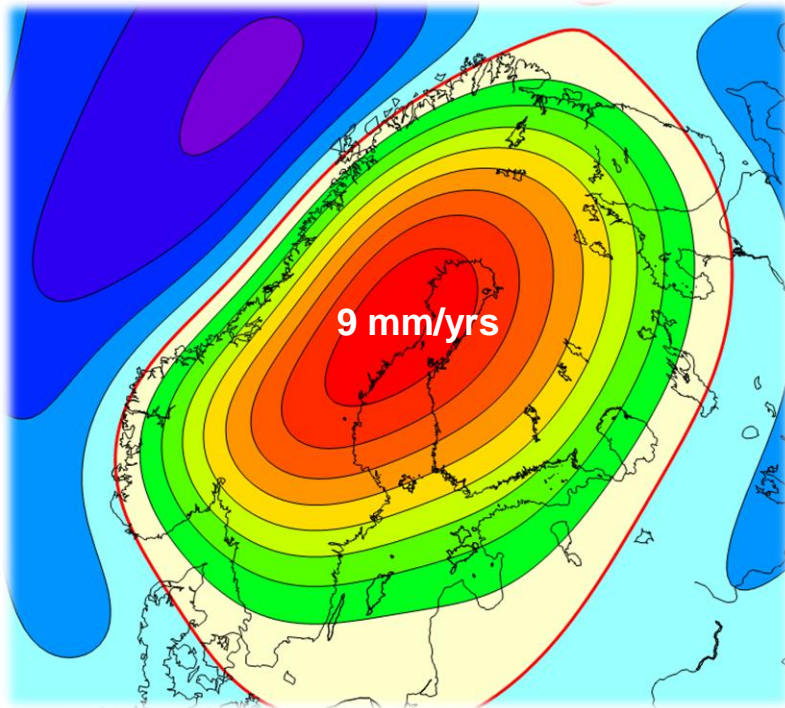
Observed palaeo shorelines



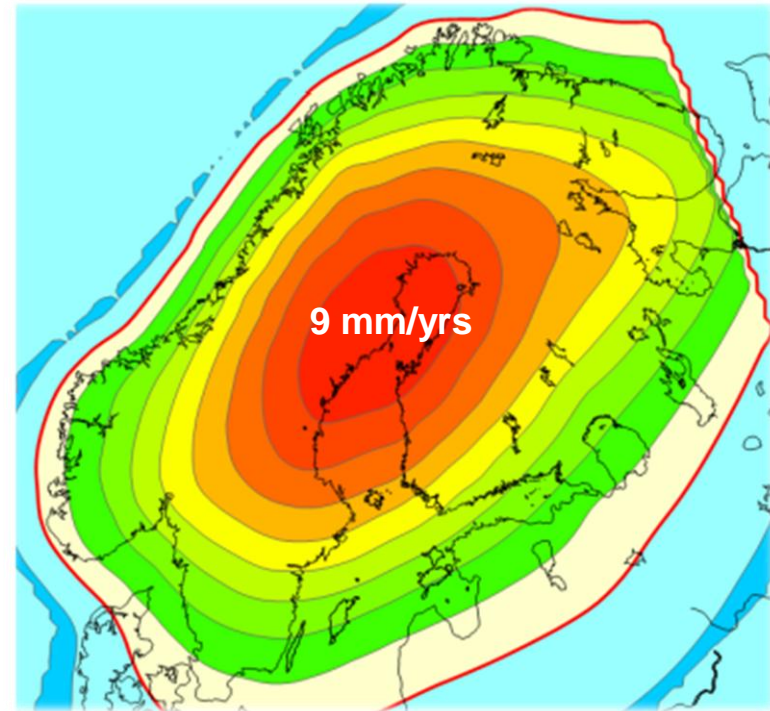
Post-glacial shorelines from Roddines, Porsangerfjord (north Norway)

Observed vs. calculated present rate of uplift

Calculated



Observed

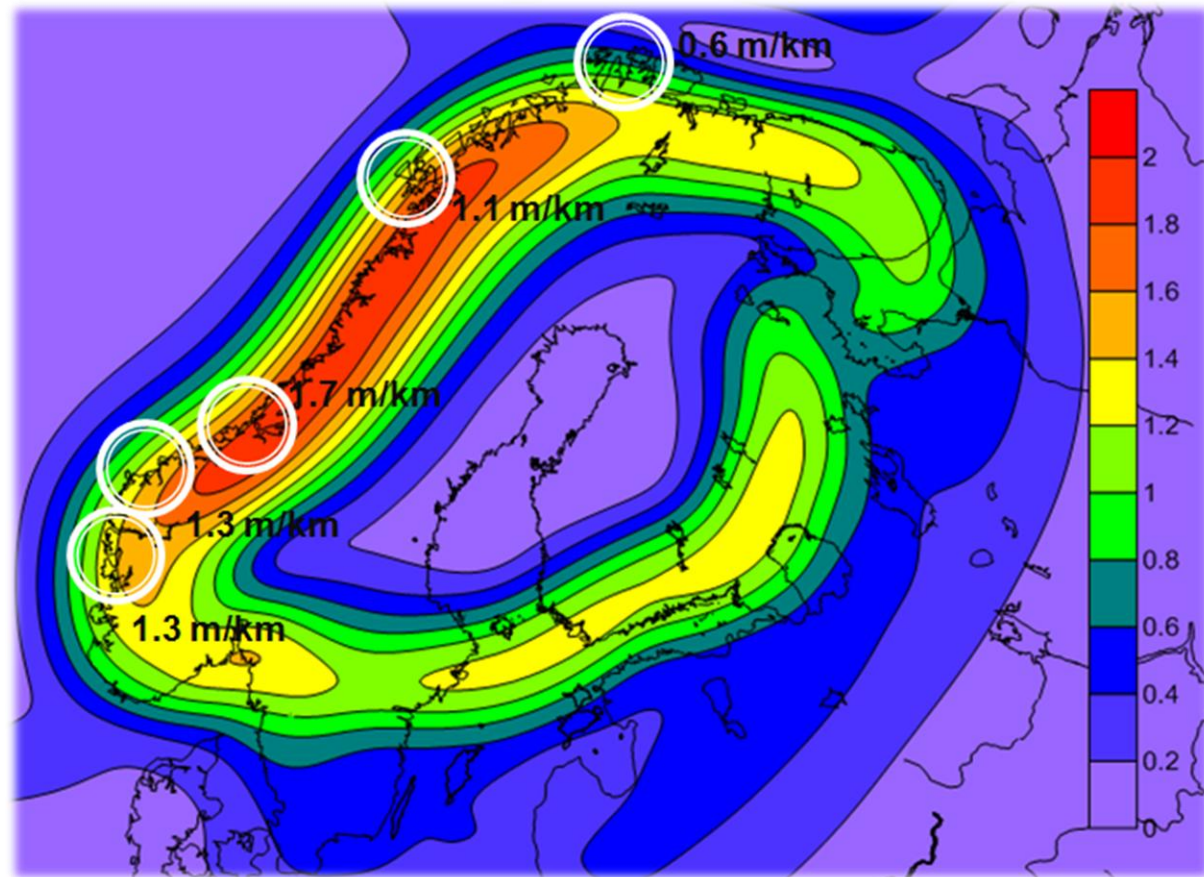


Observed (Vestøl, 2006)
- corrected for eustasy

Effect of ice model

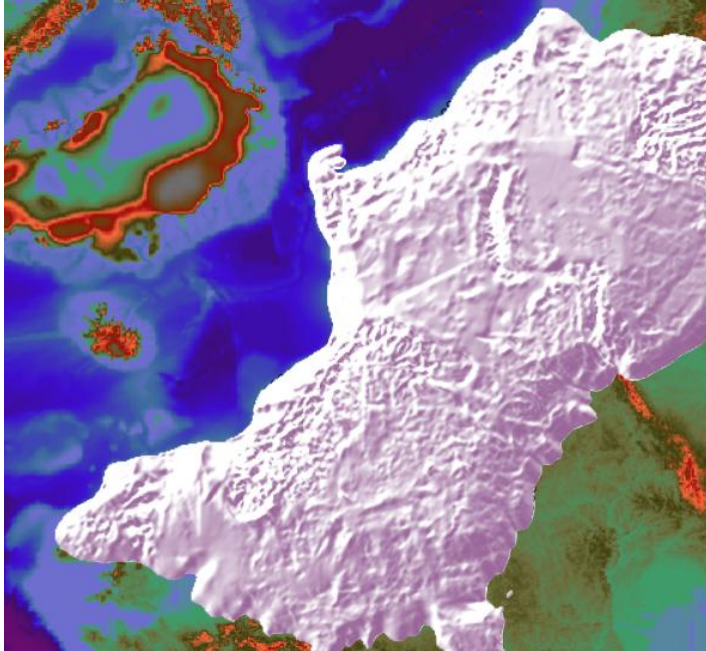
Earth model:

- A low viscosity asthenosphere (2.5×10^{19} Pa s)
- Lithosphere rigidity 5×10^{23} Nm (40 km)

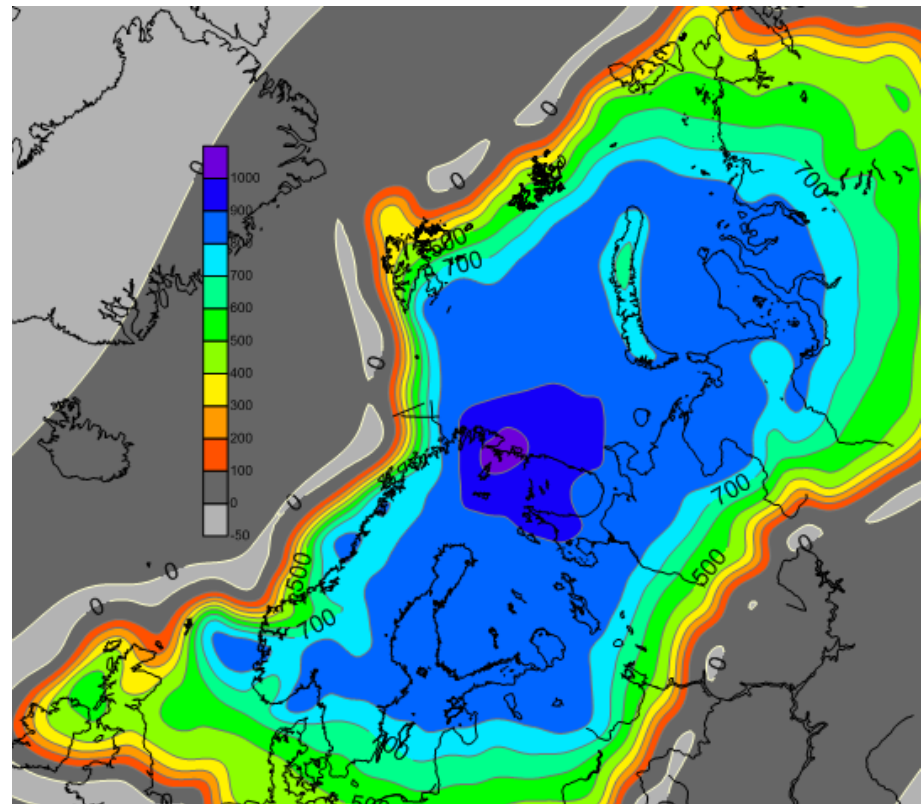


Plio-Pleistocene glacier and isostasy

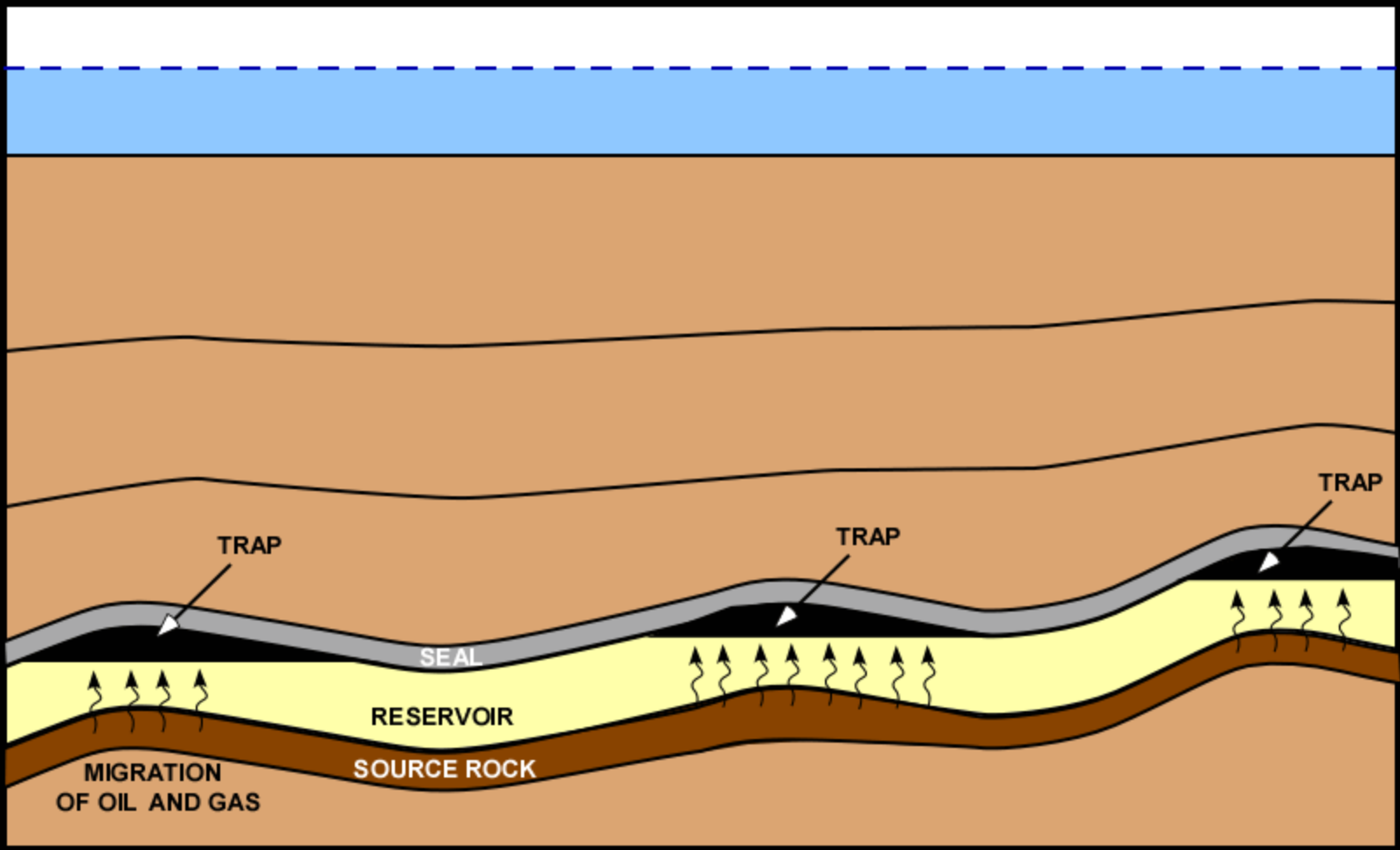
< 1.0 Ma



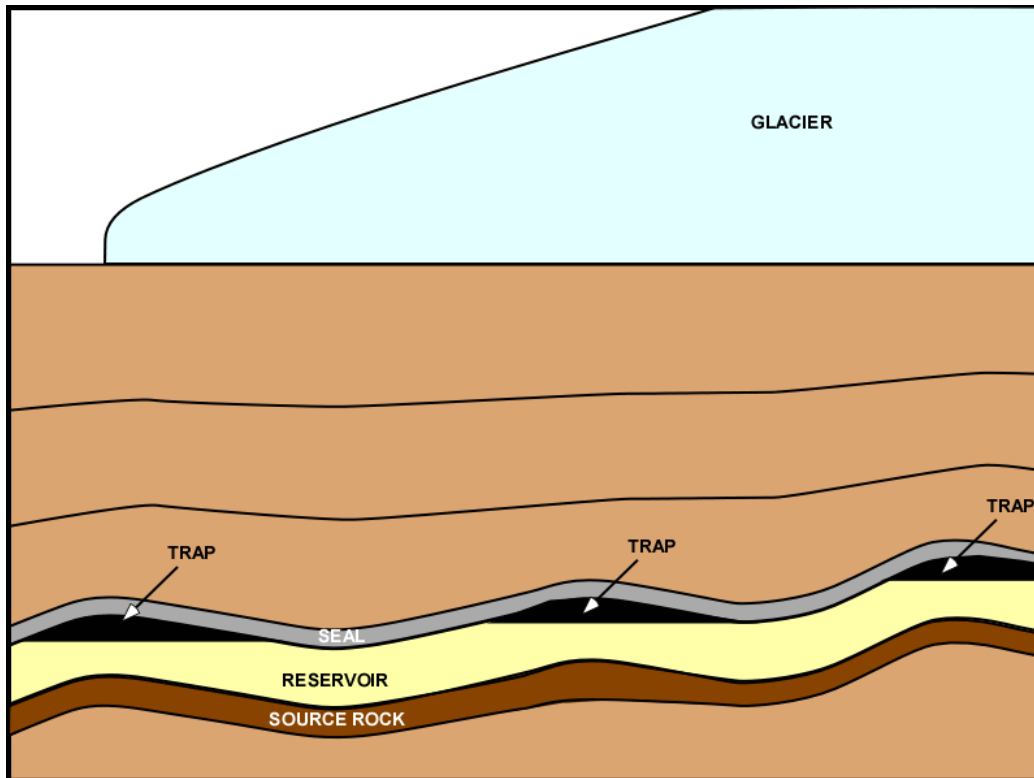
Isostasy



Petroleum system



Isostatic effects on petroleum system

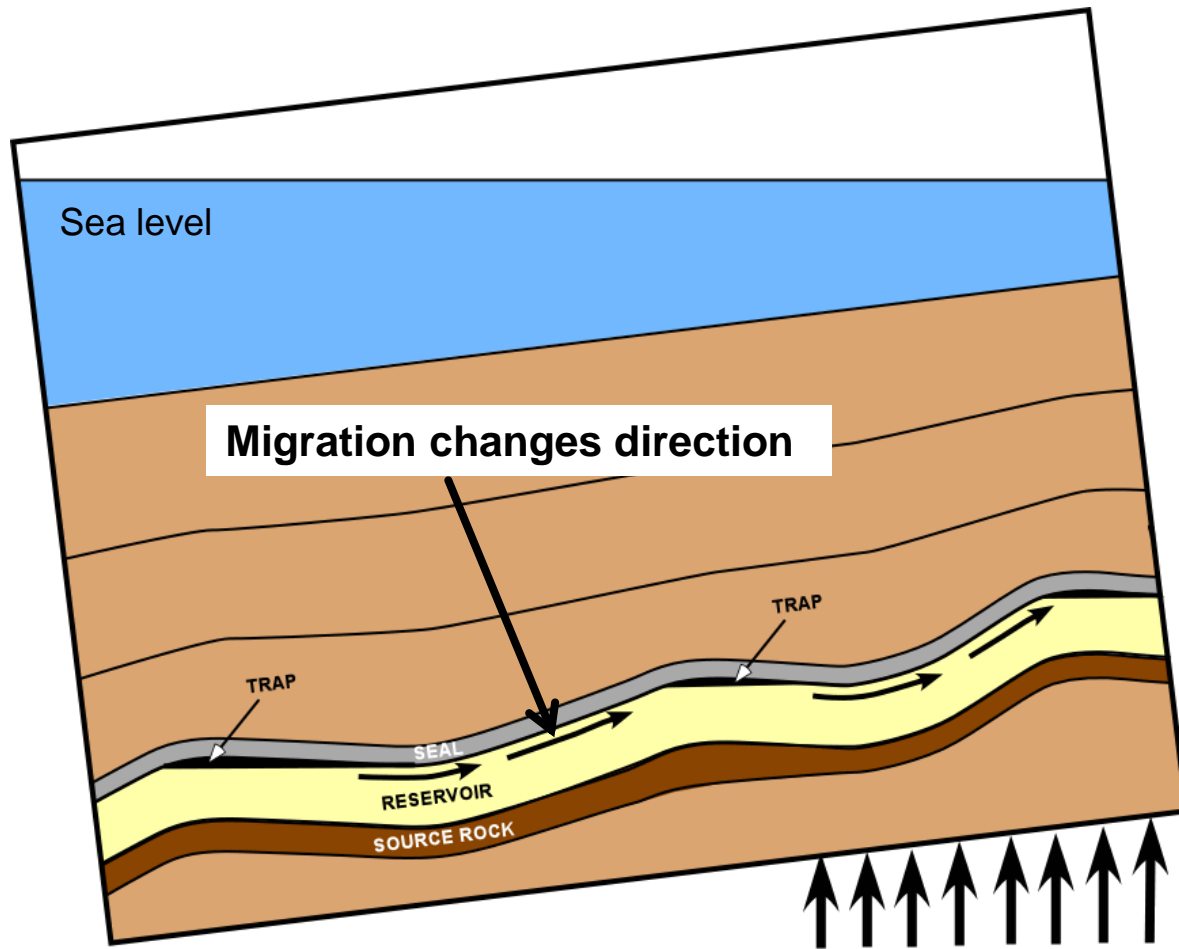


Basin in isostatic equilibrium during glaciation.

Petroleum system is located in the area near the edge of the glacier.

What happens when the ice melts?

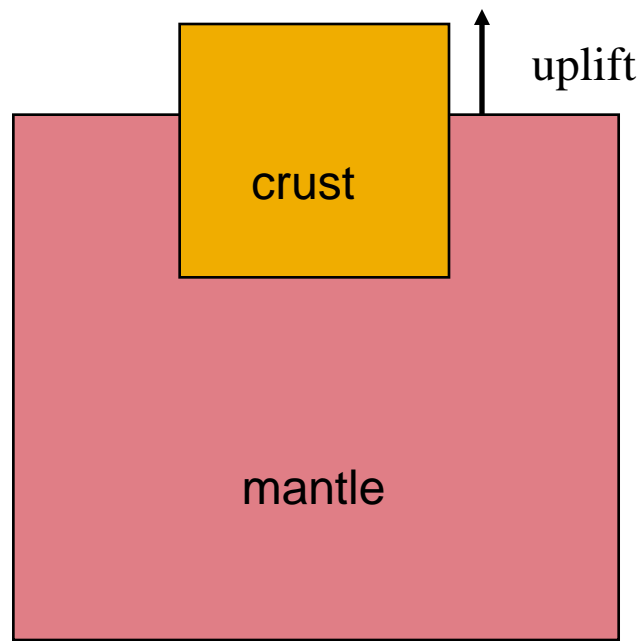
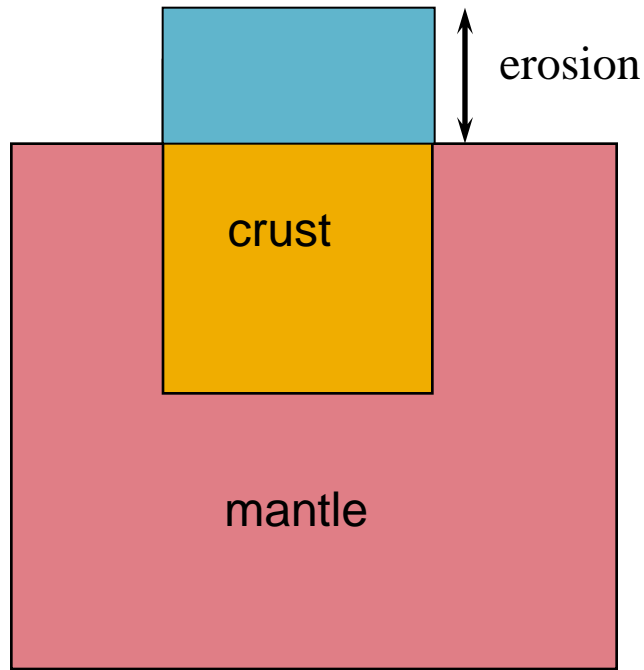
Isostatic effects on petroleum system



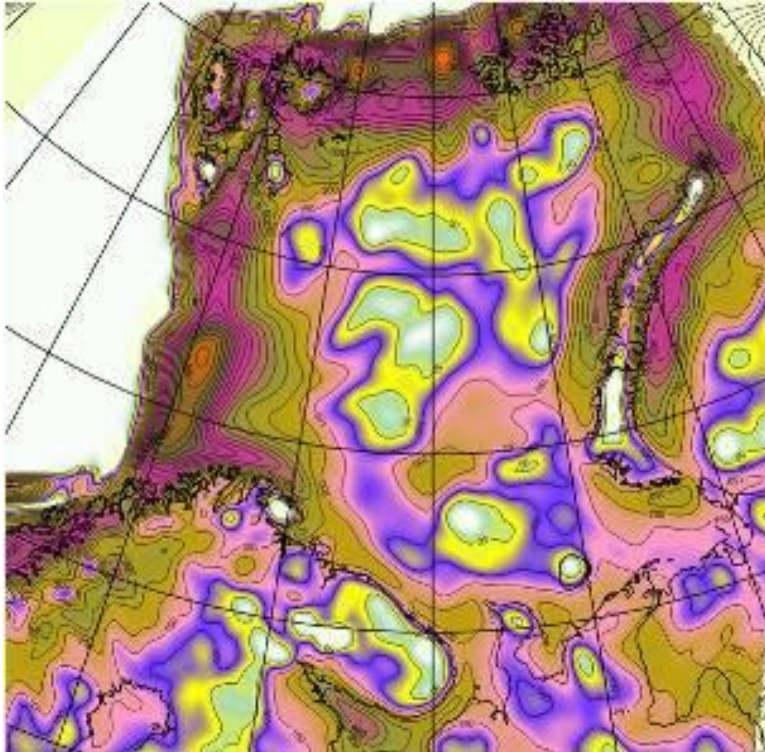
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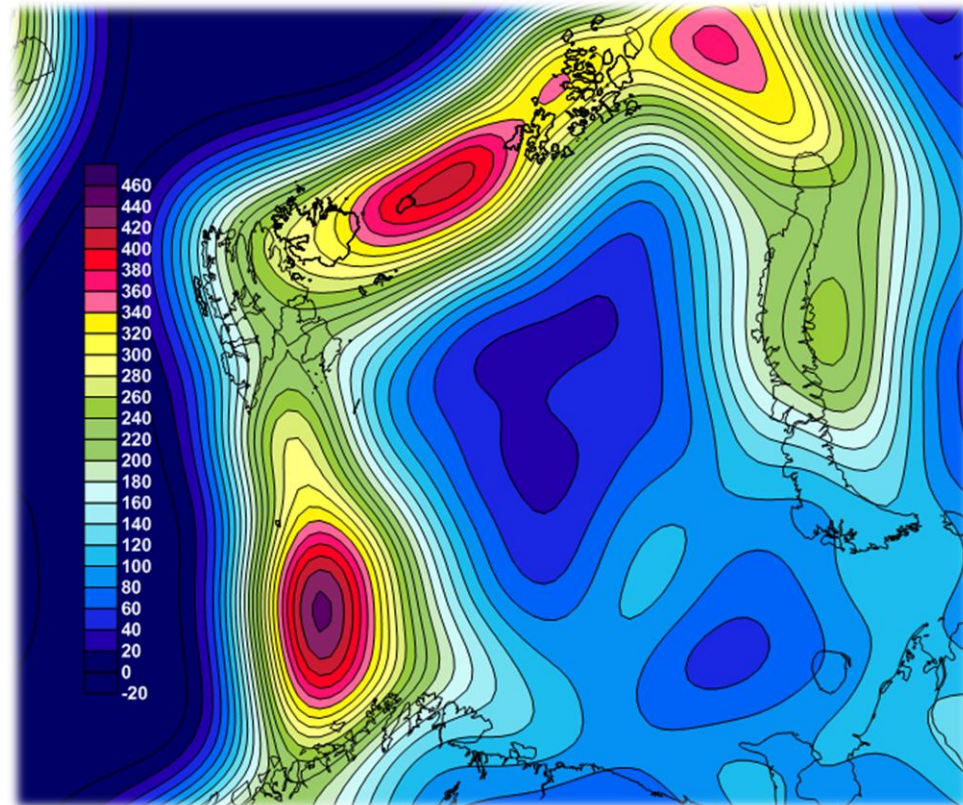
Erosion and uplift



Glacial erosion and uplift



Isostasy

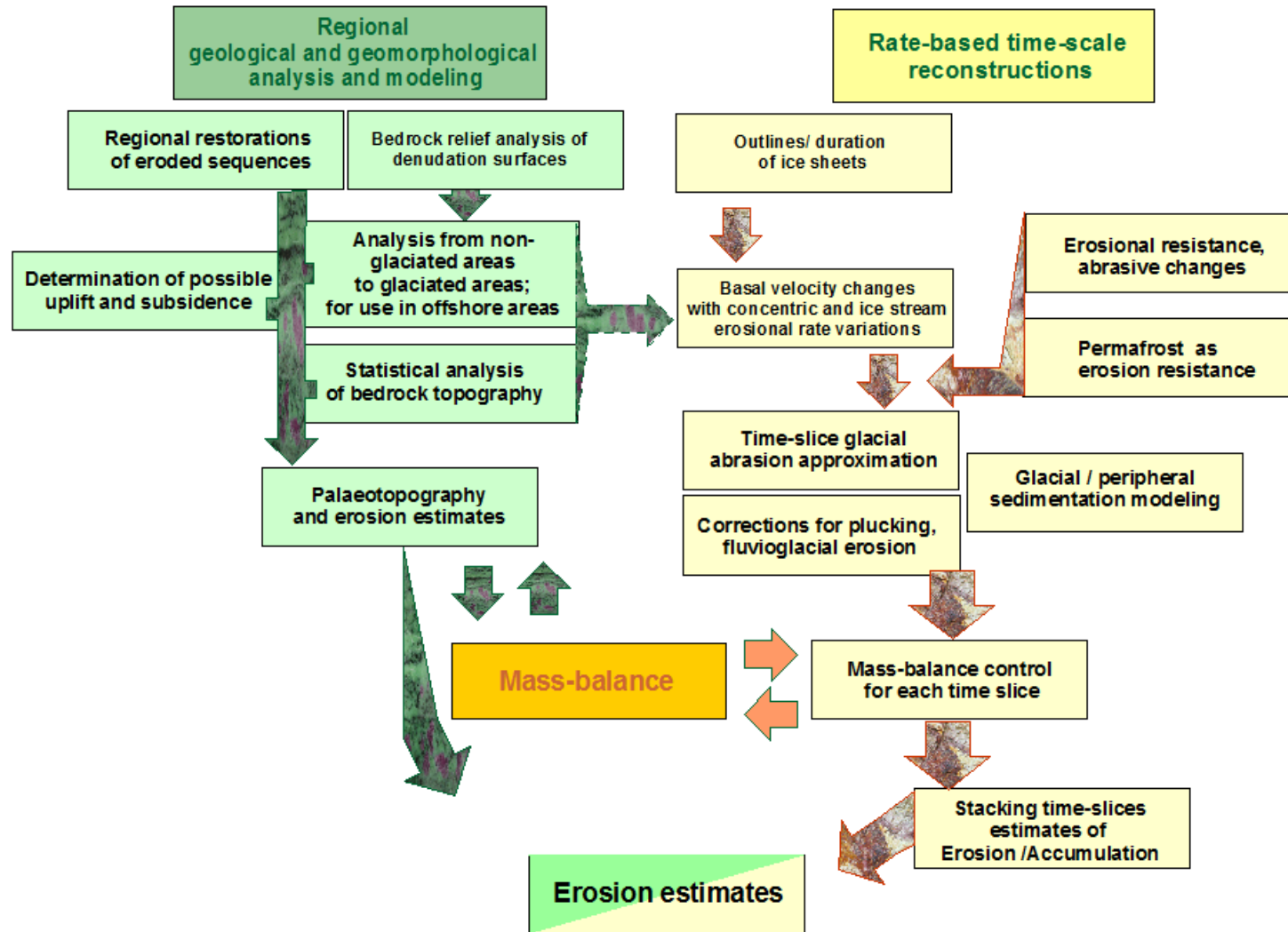


Conclusion-1



The isostatic effects of glaciers, glacial erosion/deposition are calculated. The Earth rheology (elastic lithosphere thickness and asthenosphere viscosity) is found from high resolution modelling of the rebound after the last glaciation. Based on this rheology it is shown that the tilting of the reservoirs in the western Barents Sea could be significant, up to 2m/km

Neogene and Paleogene erosion



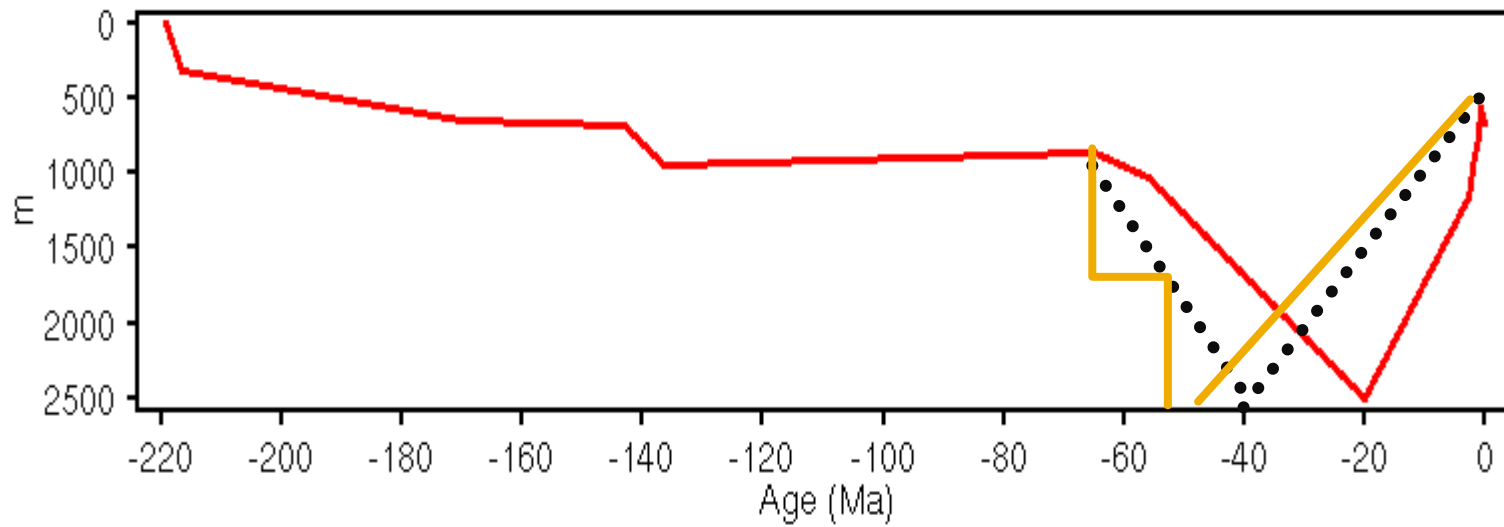
Observed vitrinite reflectance

Time of max burial

Stratigraphy within missing section

Thermal conductivities of the missing section

Timing of Max burial?

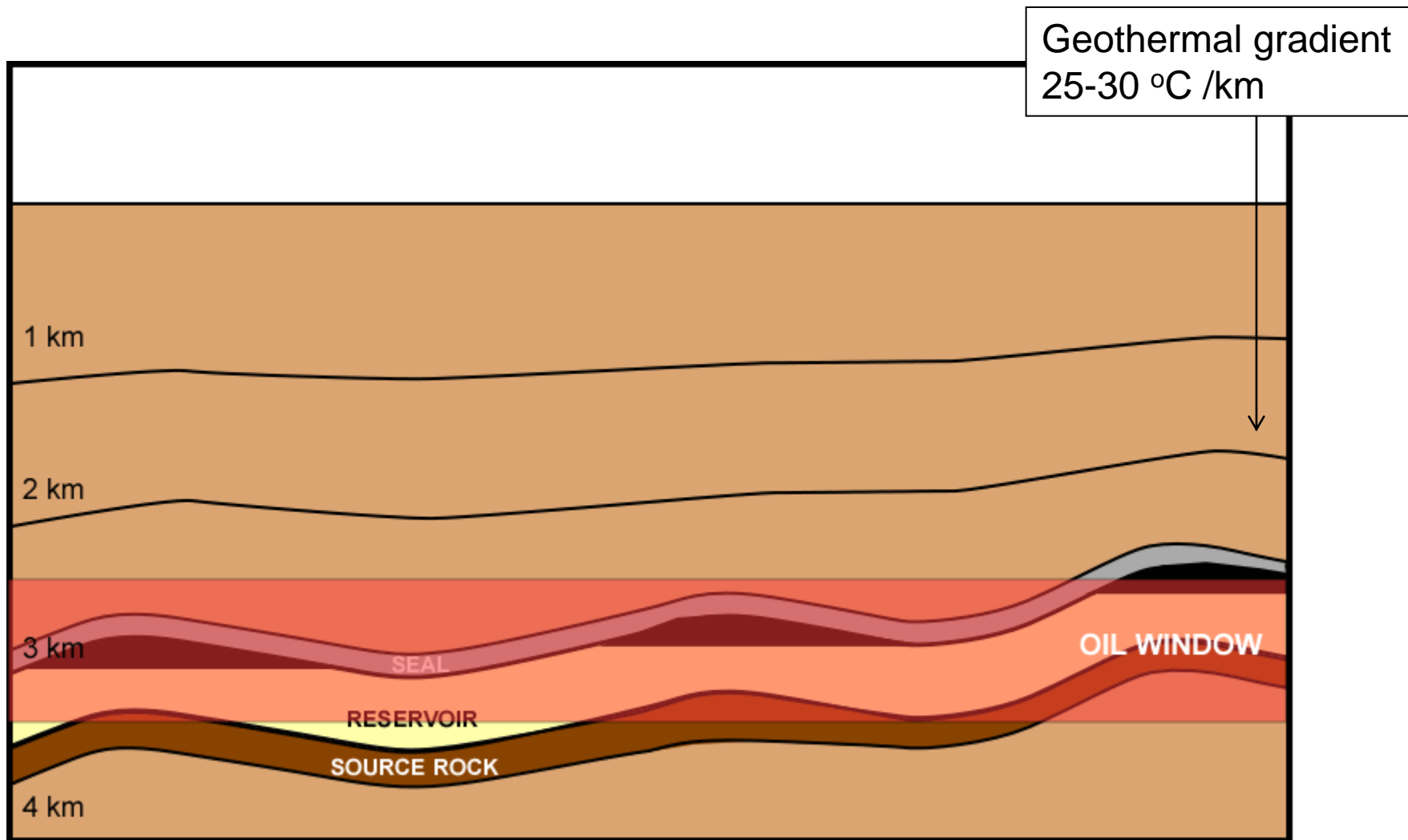


Modelling of 2D transects

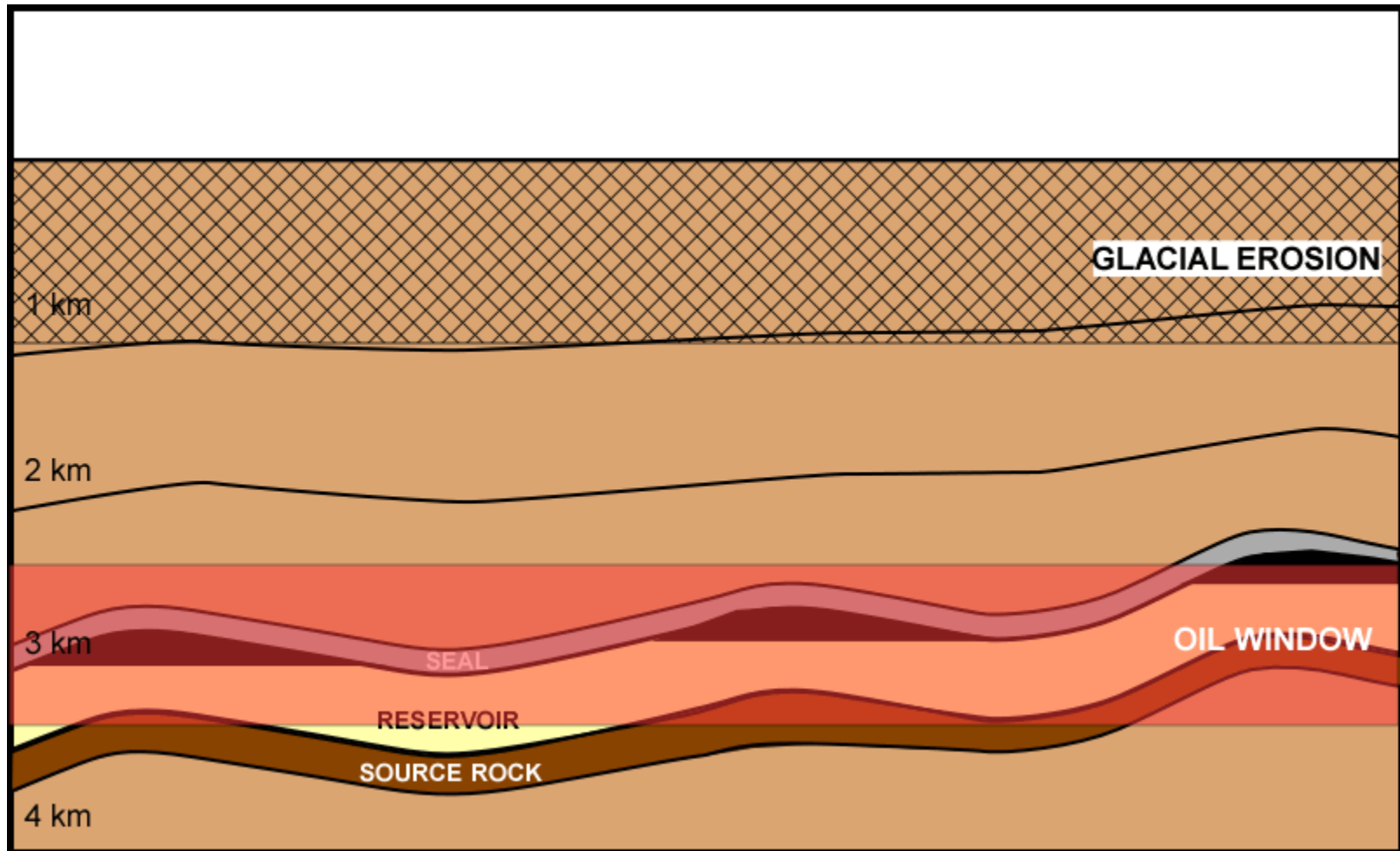
Aim:

**temperature and maturity effects
of Plio- Pleistocene glaciations
and Neogene erosion**

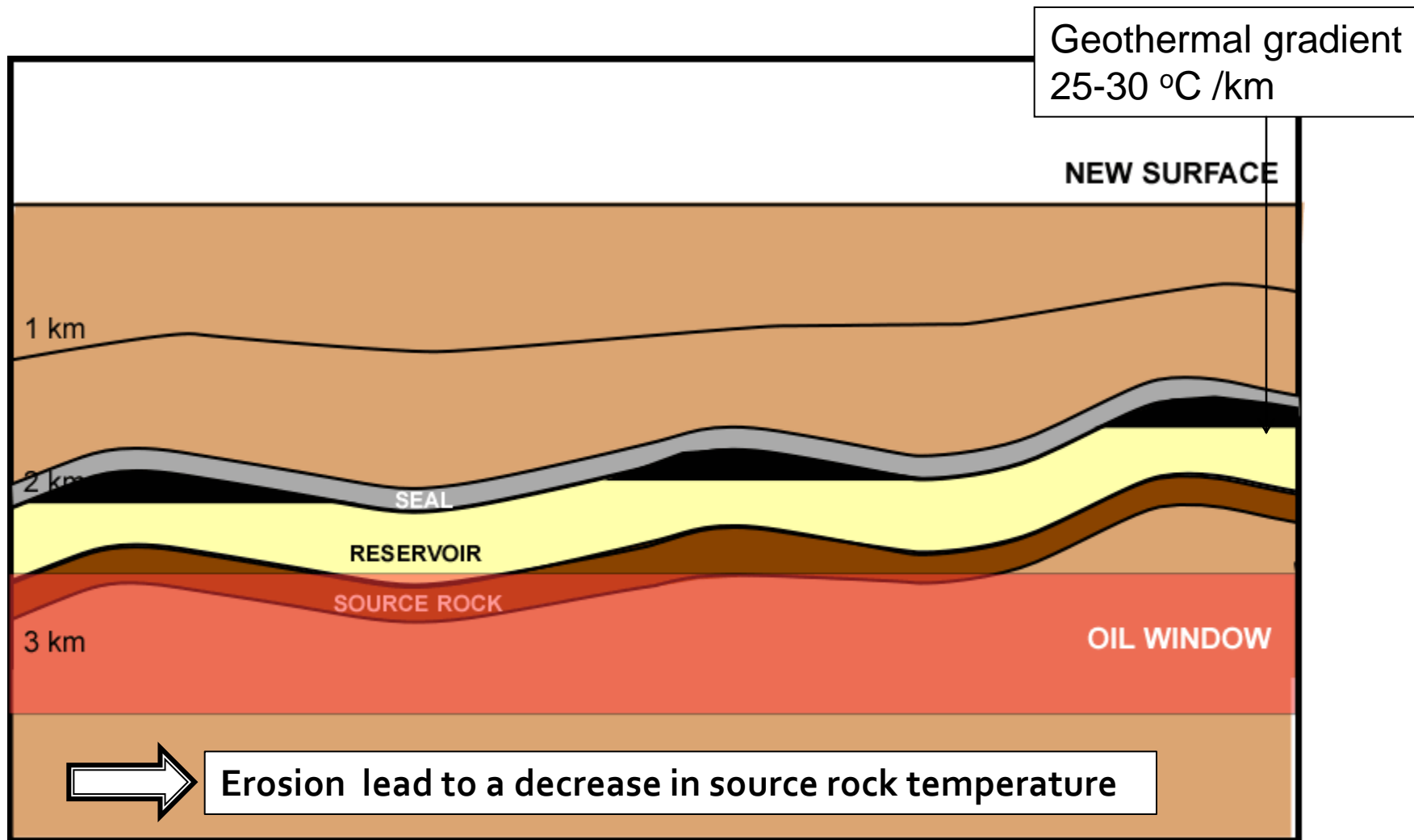
Effects of erosion on petroleum systems



Effects of erosion on petroleum systems



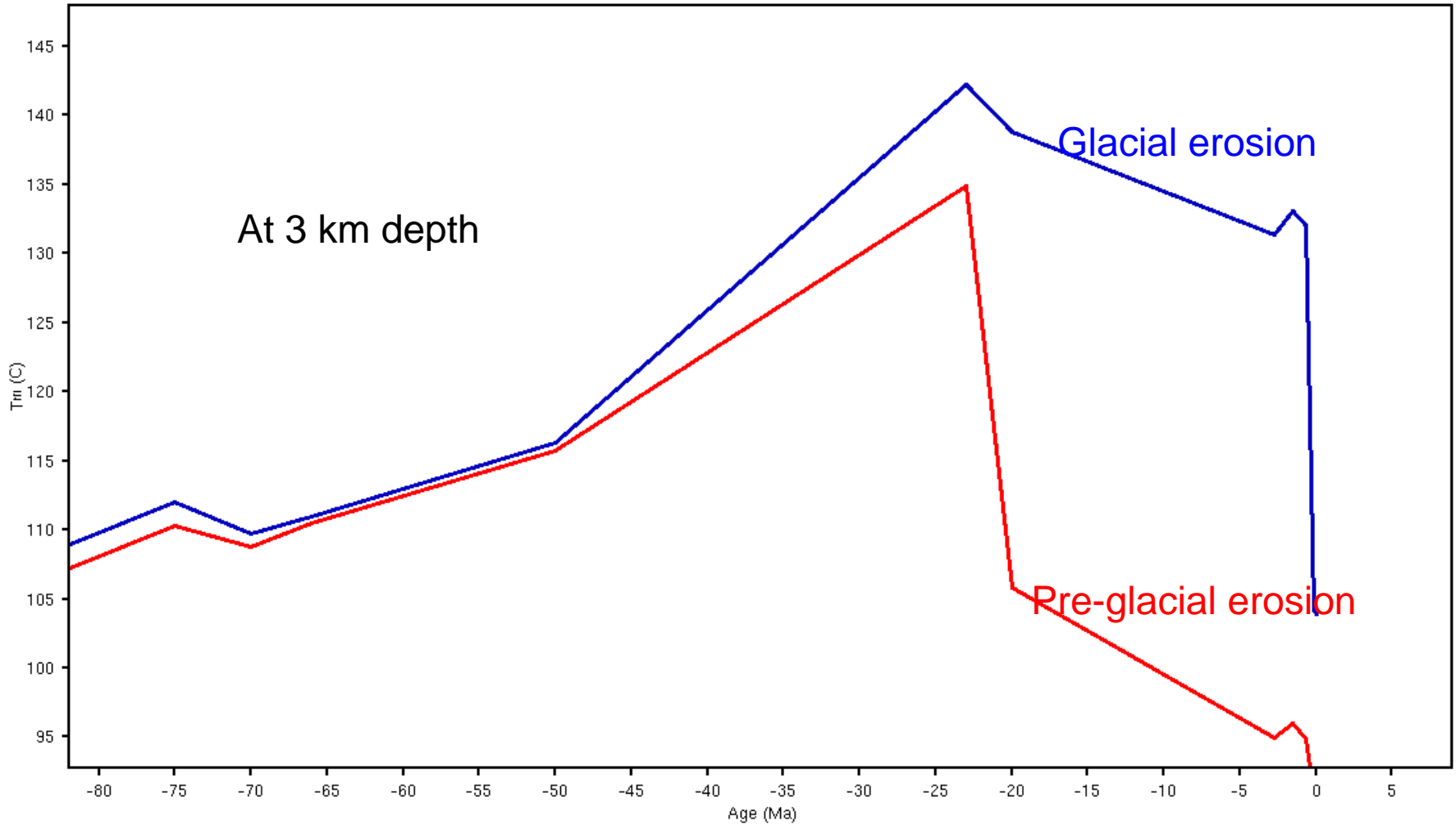
Effects of erosion on petroleum systems



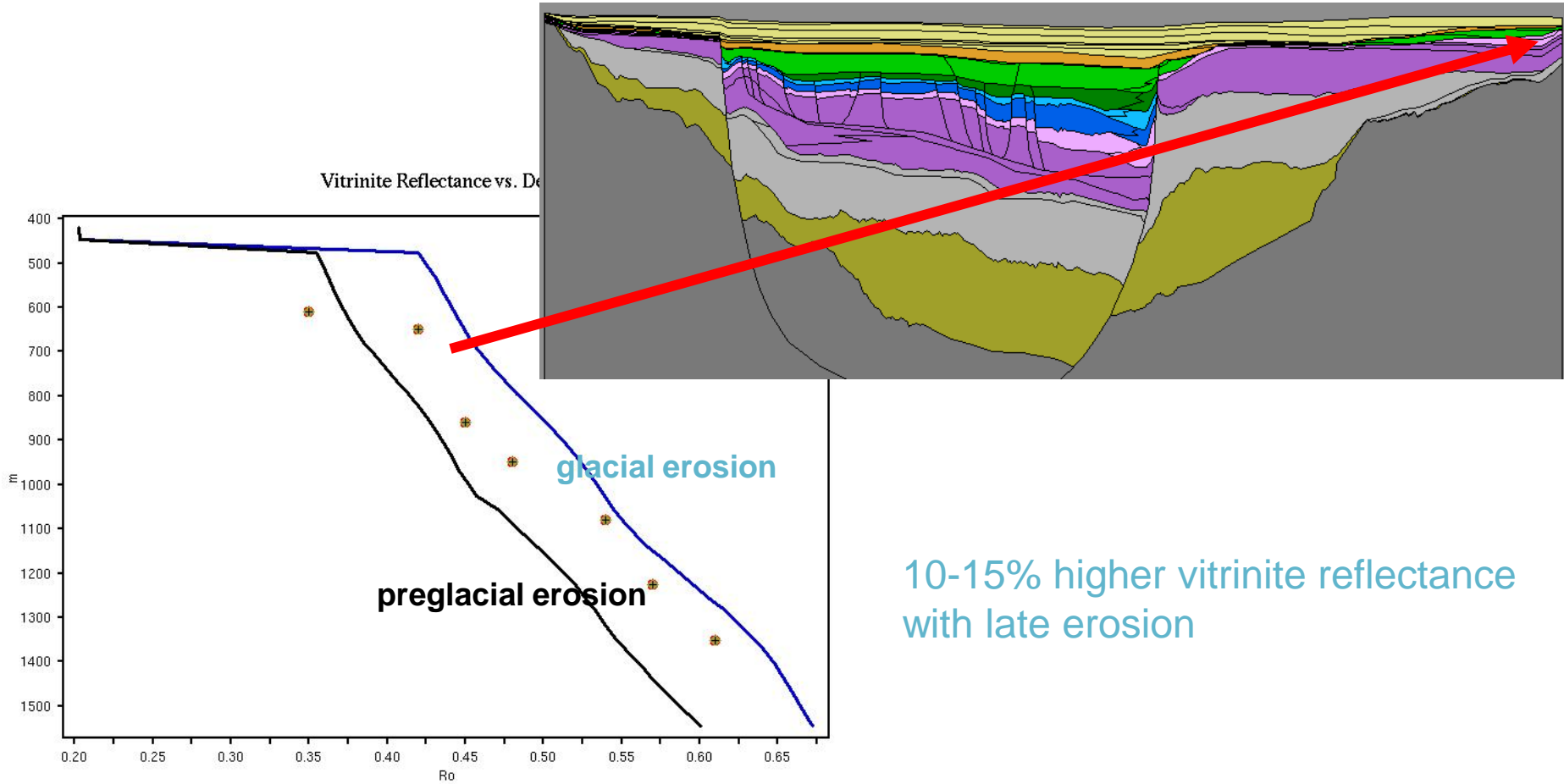
Temperature effect of pre-glacial vs glacial erosion



Cell Temperature History



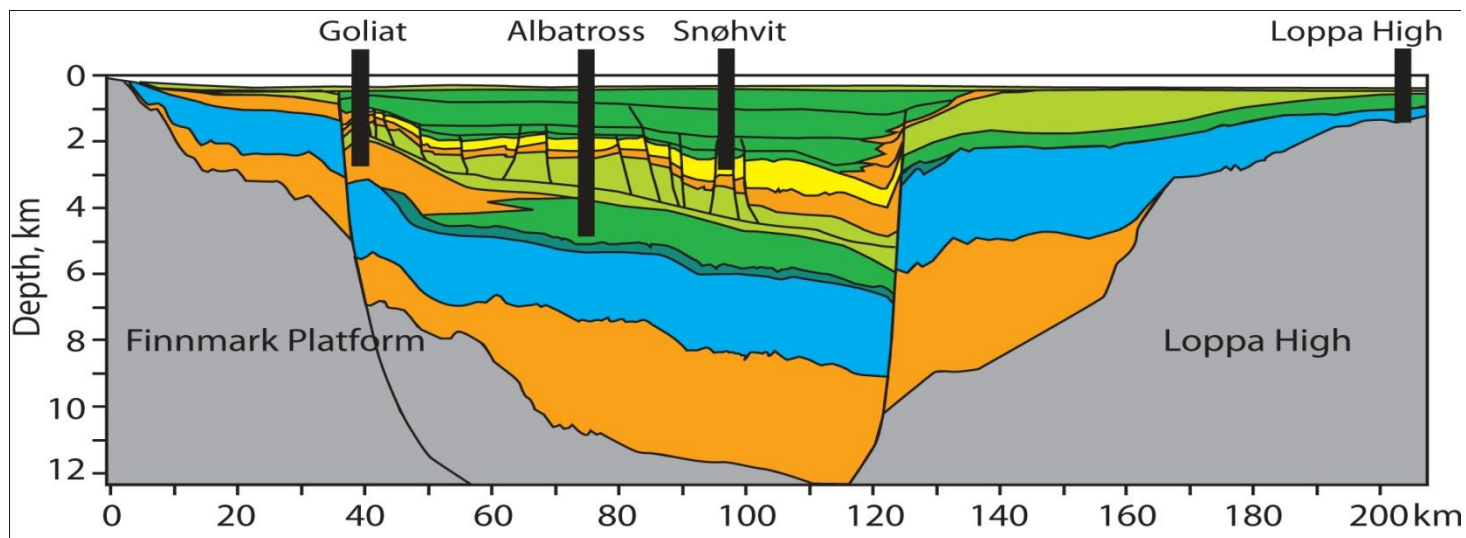
Vitrinite effect of pre-glacial vs glacial erosion



Pre-glacial versus glacial erosion

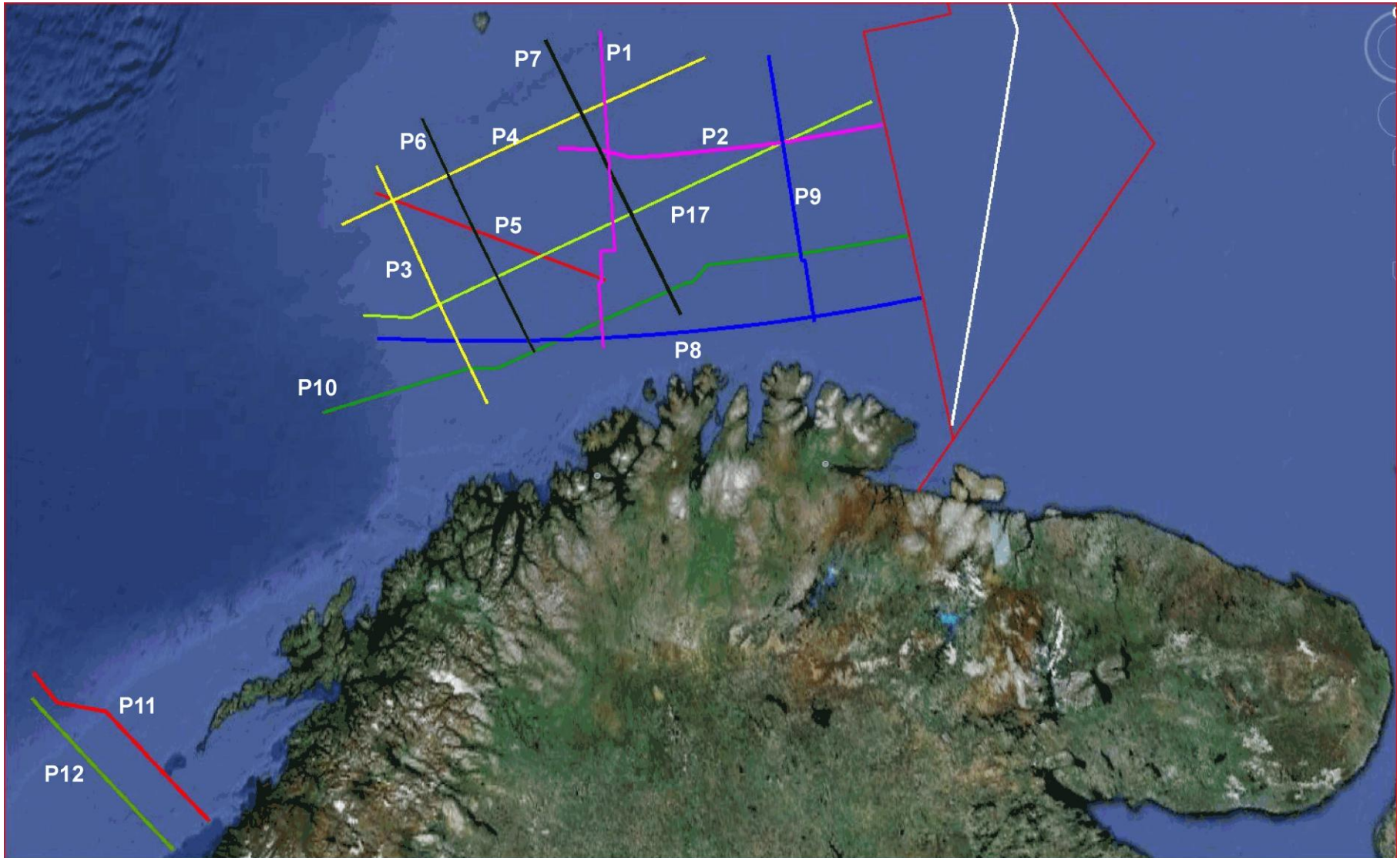
Prediction of pre-glacial erosion

period	erosion			
2.7- 1.5Ma	200m	250m	300m	350m
1.5- 0.7Ma	300m	350m	350m	400m
0.7- 0Ma	100m	100m	100m	300m



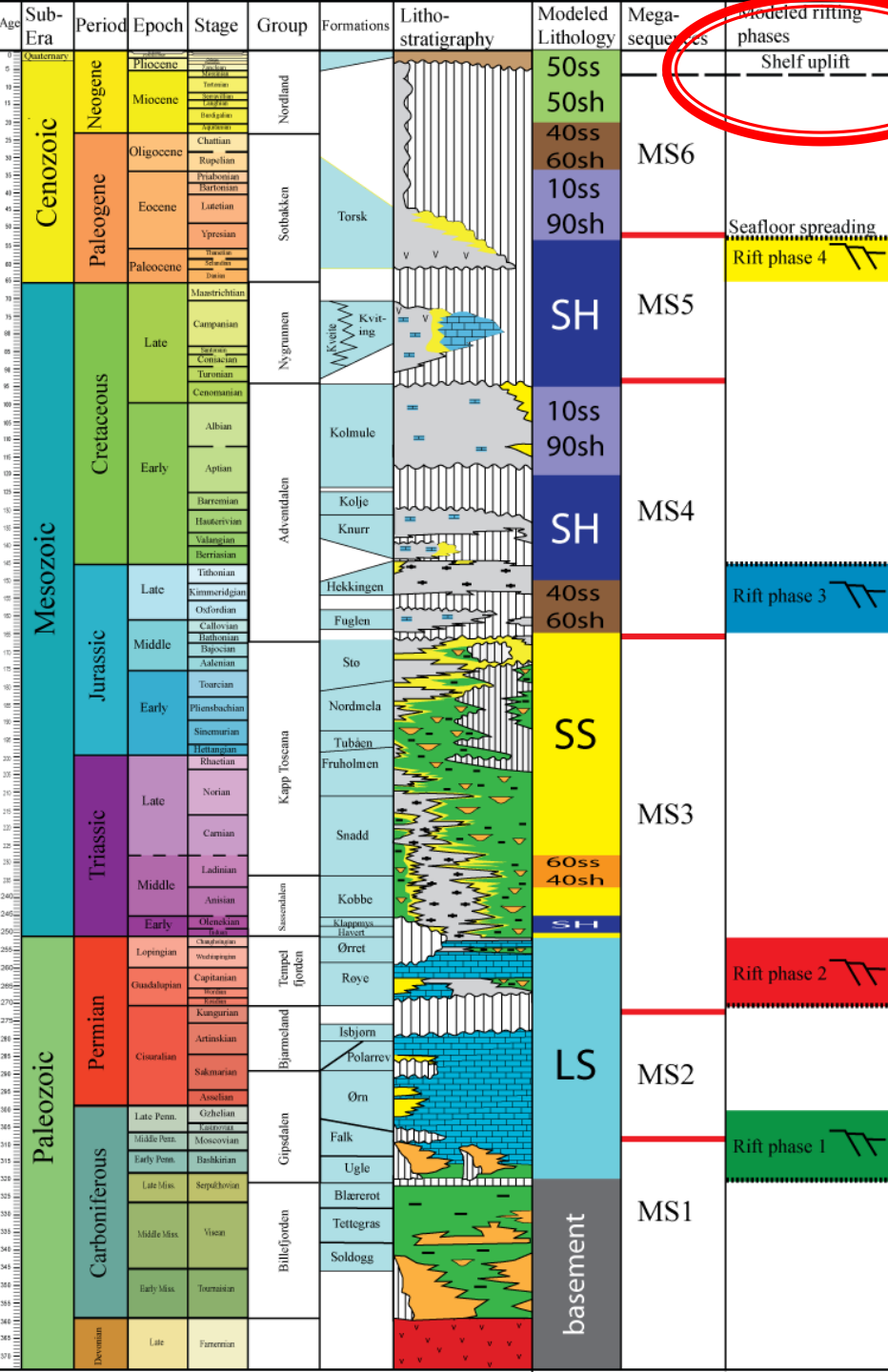
Pre-glacial	100m	150m	150m	300m
Total	700m	850m	900m	1350m

2D lines



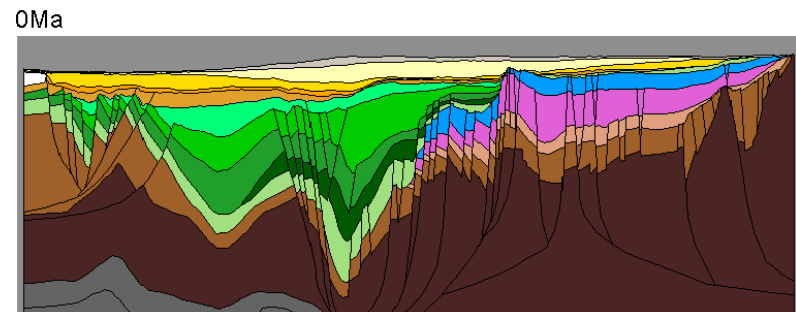


TECTONOR



Lithostratigraphic column southwest Barents Sea (from Clark et al., in review)

- An advanced 2D basin modelling system :
 - Reconstruction of the basin evolution
 - Fault restoration
 - Chemical compaction
 - Isostatic deflections (with flexure)
 - Lithospheric thinning (with necking)
 - Magmatic intrusions
 - User guided salt movements
 - Temperature/maturity effects



Vitrinite reflectance gives us the opportunity to constrain the Neogene erosion